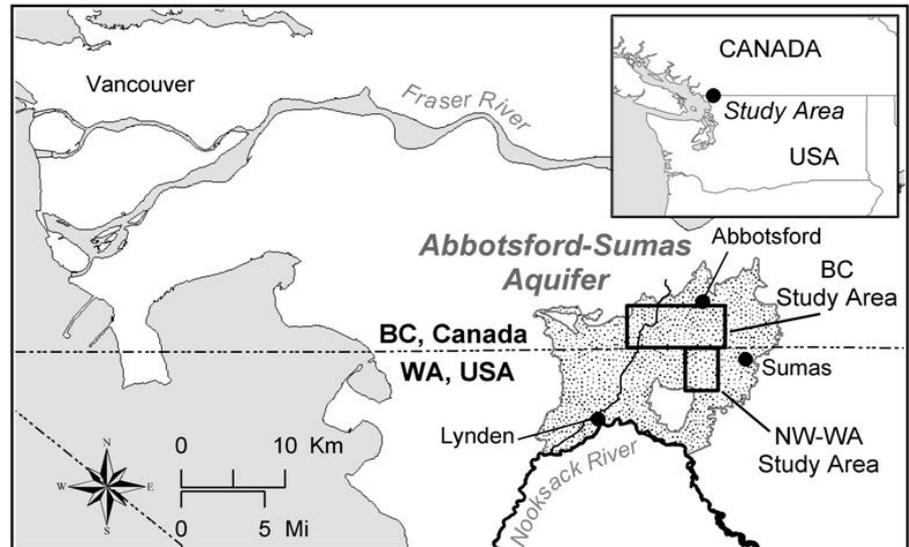


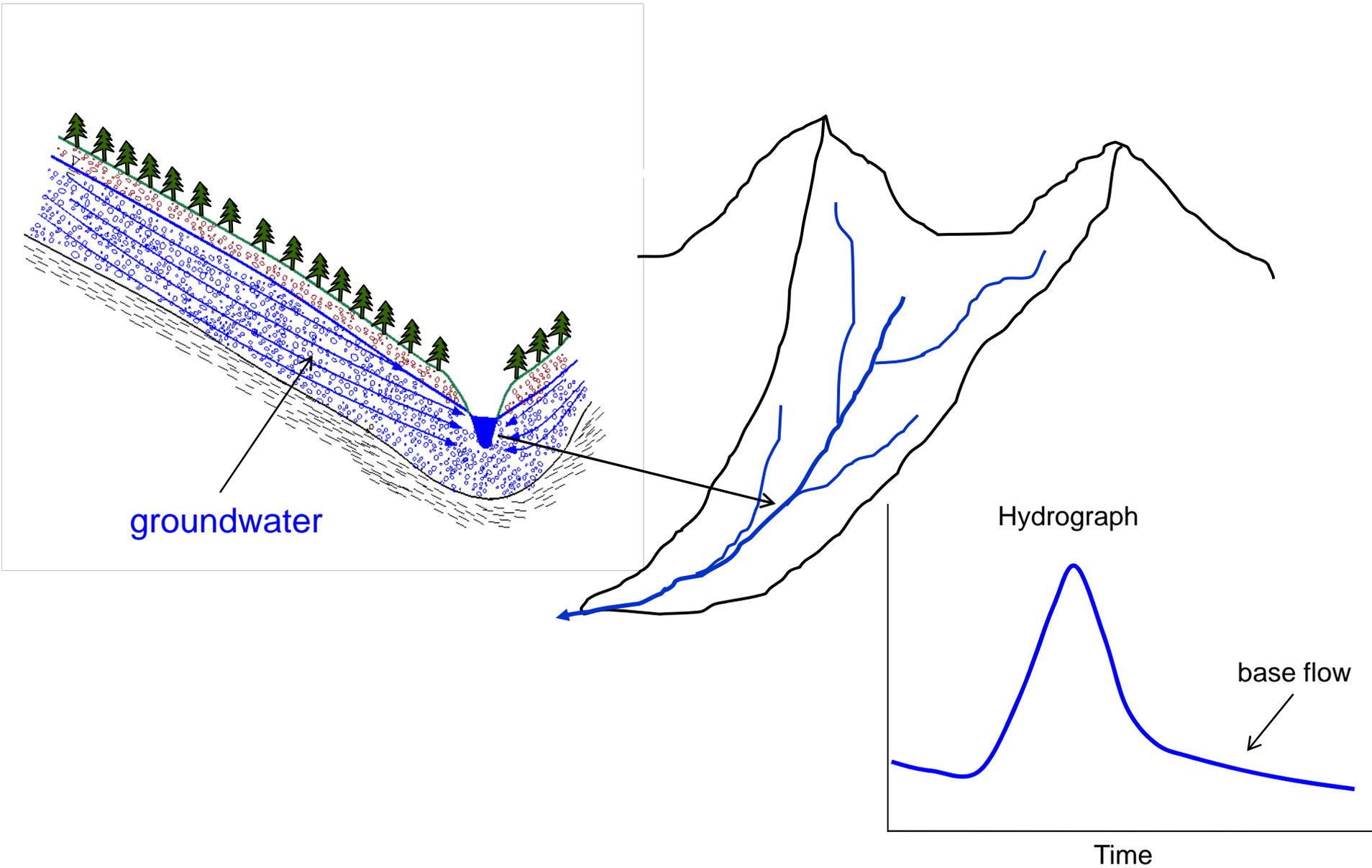
# Lecture 21: Groundwater: Hydraulic Conductivity

## Key Questions

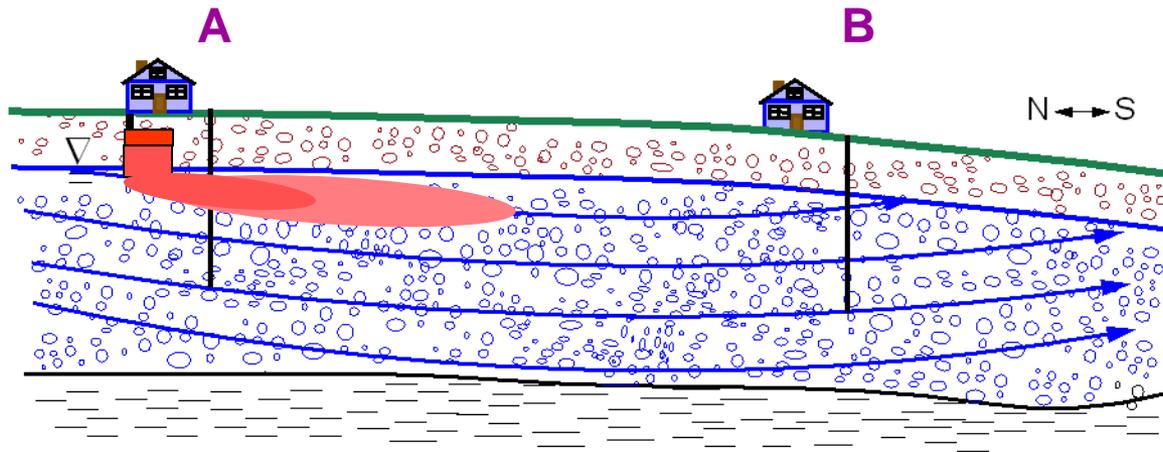
1. What causes groundwater to move?
2. What is the hydraulic conductivity?
3. What is Darcy's Law?
4. How is groundwater velocity estimated?
5. What makes good aquifers in the Puget Lowlands?
6. What is a cone of depression?



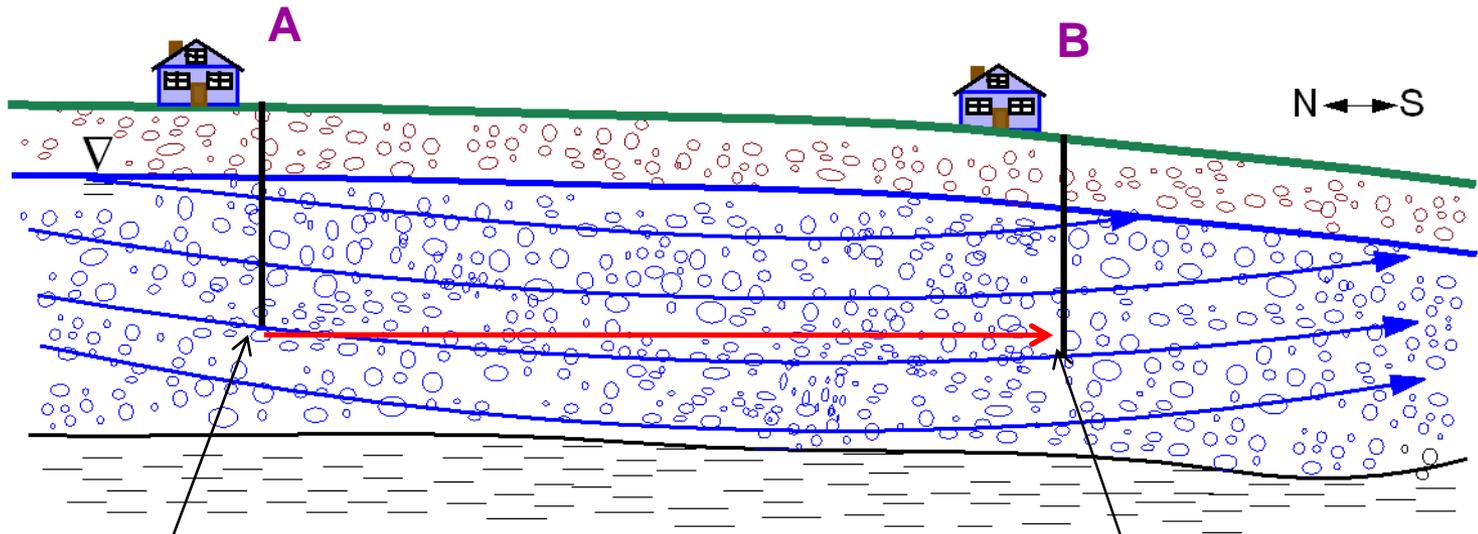
Groundwater supports streamflow in between rain events (baseflow)



# Groundwater movement controls the transport of contaminants



A hydraulic gradient causes water to move

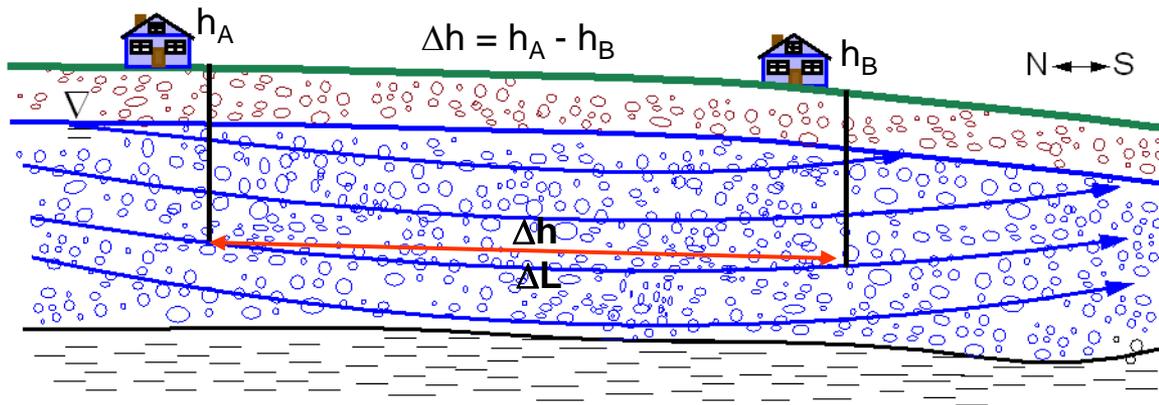


water has hydraulic head (pressure and gravitational energy) at location A

water has hydraulic head (pressure and gravitational energy) at location B

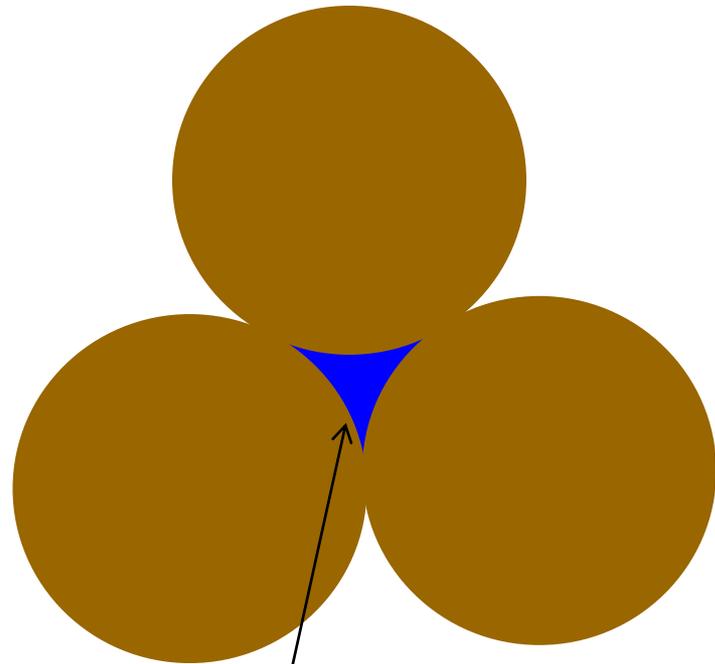
The **hydraulic gradient** is the change in total head divided the distance over which the change occurs.

$$\text{hydraulic gradient} = \Delta h / \Delta L$$



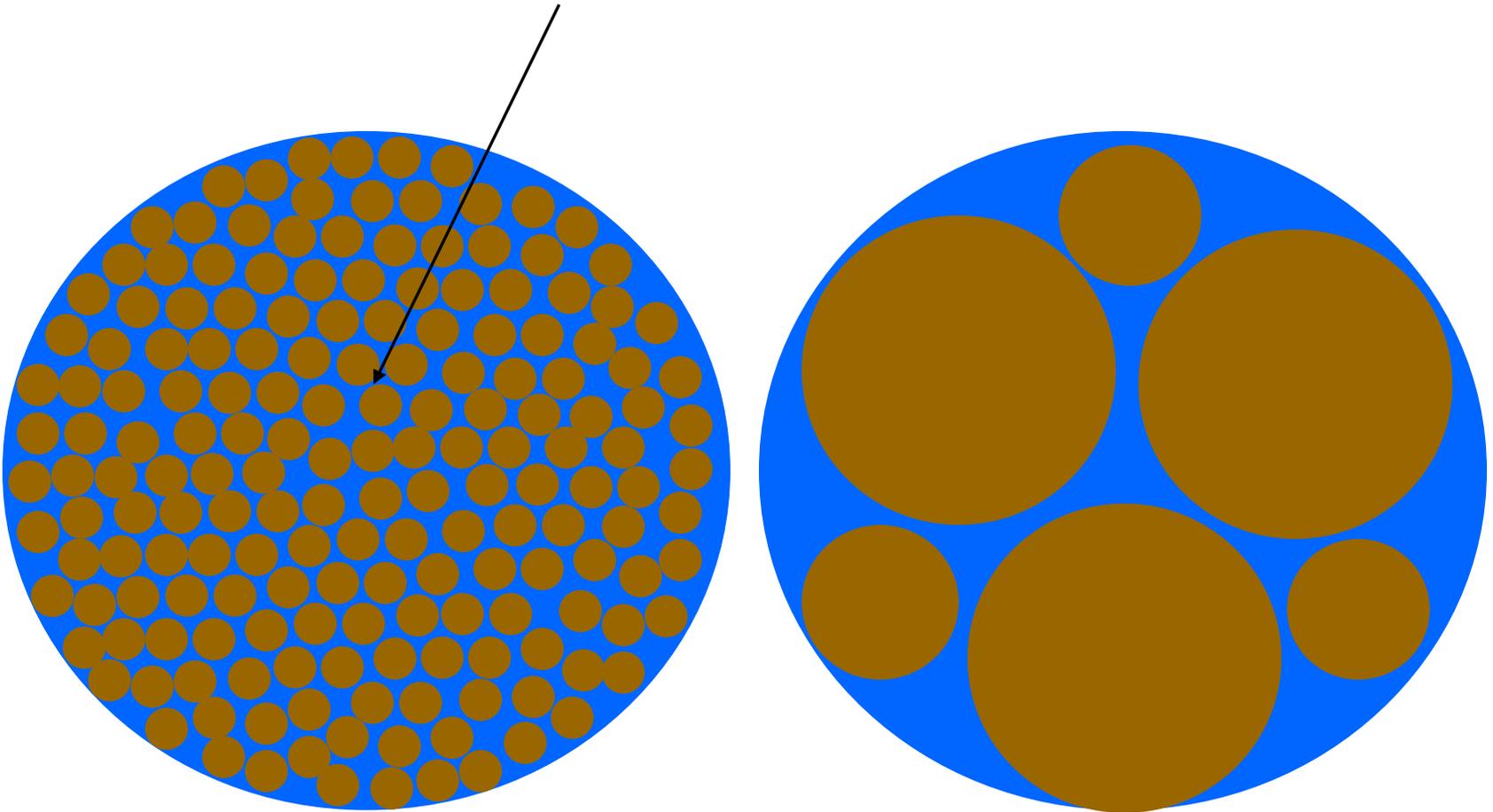
The hydraulic conductivity of the sediment will resist the water flow

small area available for flow,  
low hydraulic conductivity



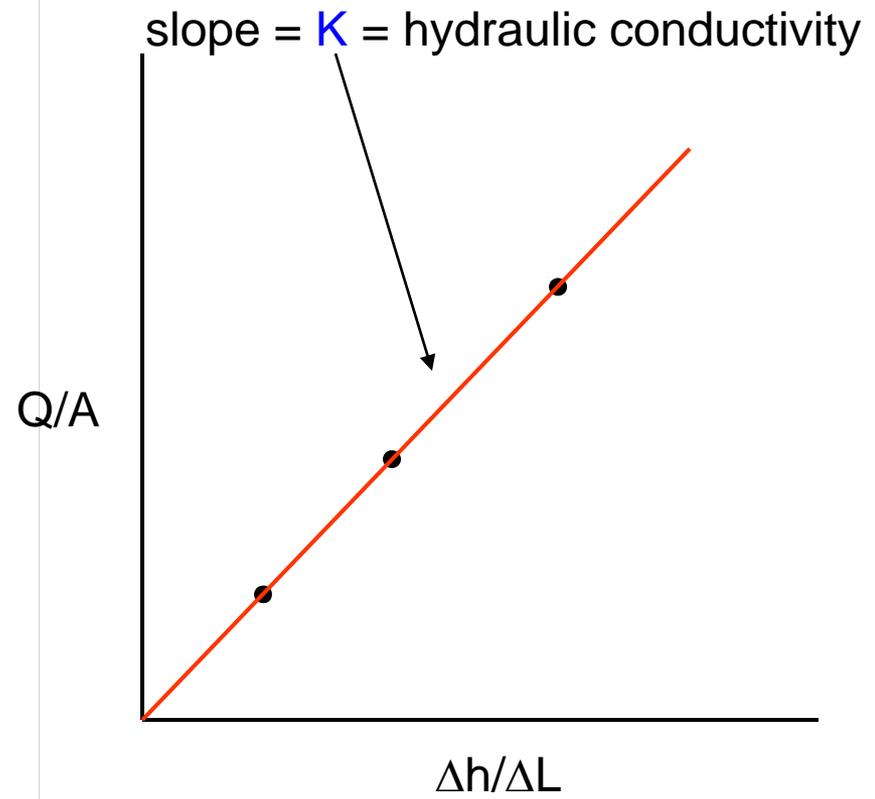
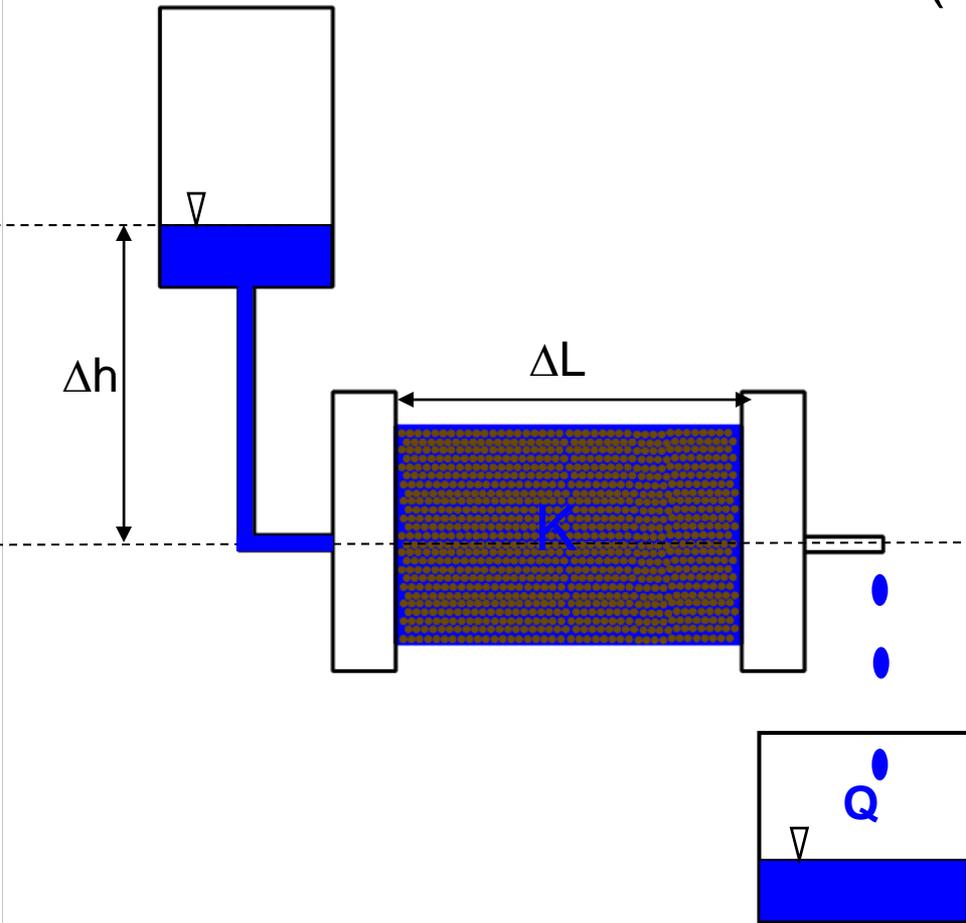
large grains, large area available for flow, large hydraulic conductivity

Smaller grains, means smaller pores (area available for flow), and more frictional resistance, and hence lower hydraulic conductivity

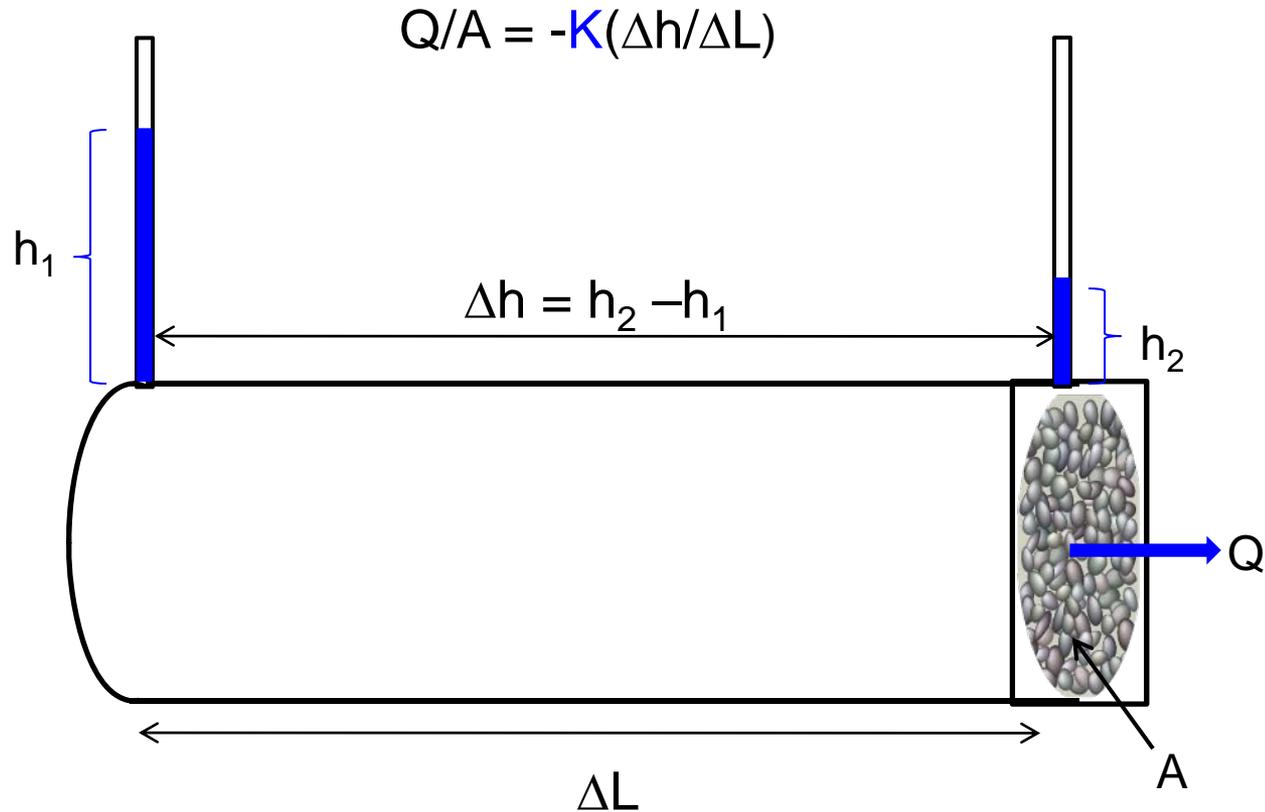


# Darcy's Law

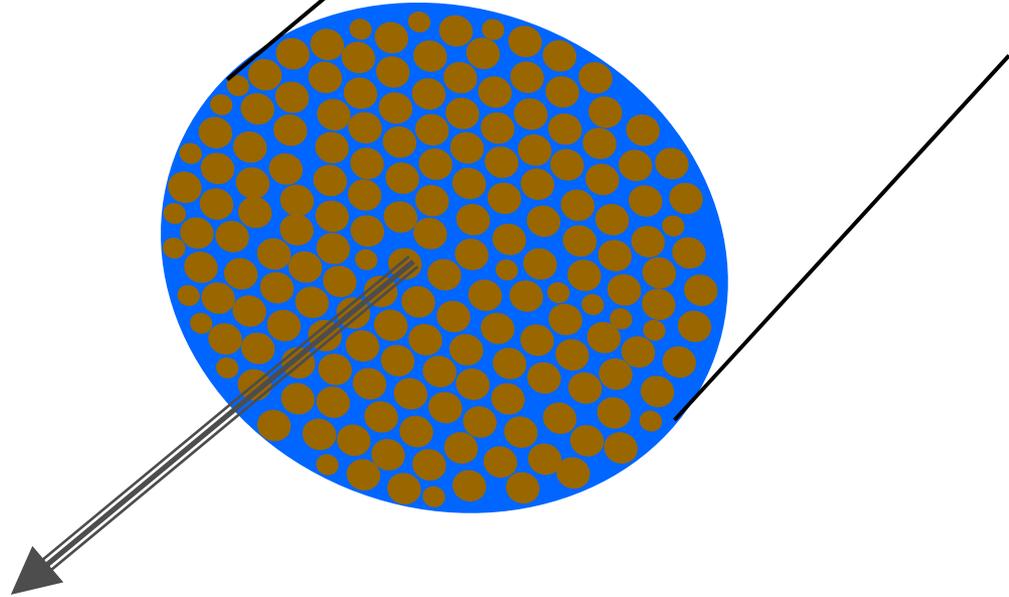
$$Q/A = -K(\Delta h/\Delta L)$$



Darcy's Law states that the amount of water ( $Q$ ) flowing through porous media depends on the energy driving the water flow ( $\Delta h/\Delta L$ ) and the hydraulic conductivity ( $K$ ) of the porous media

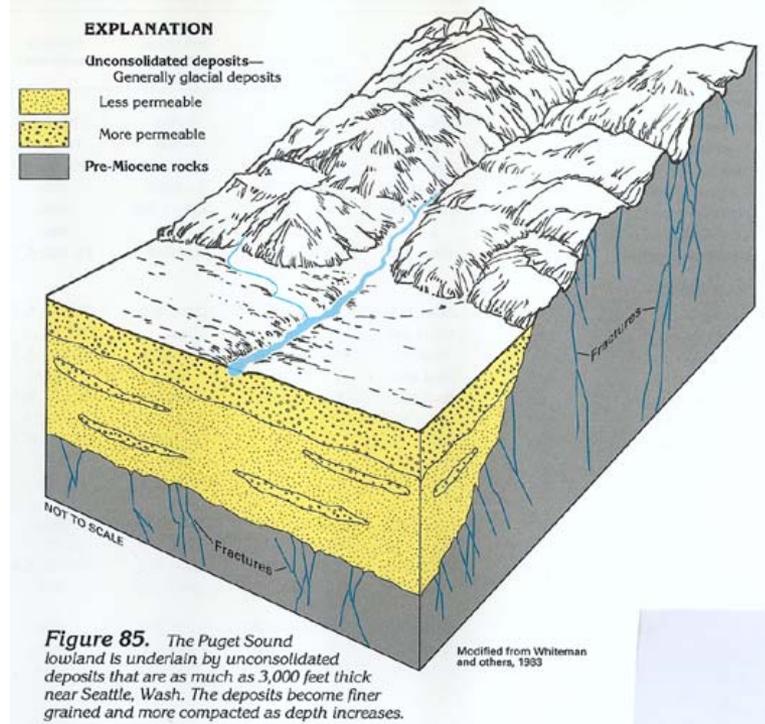
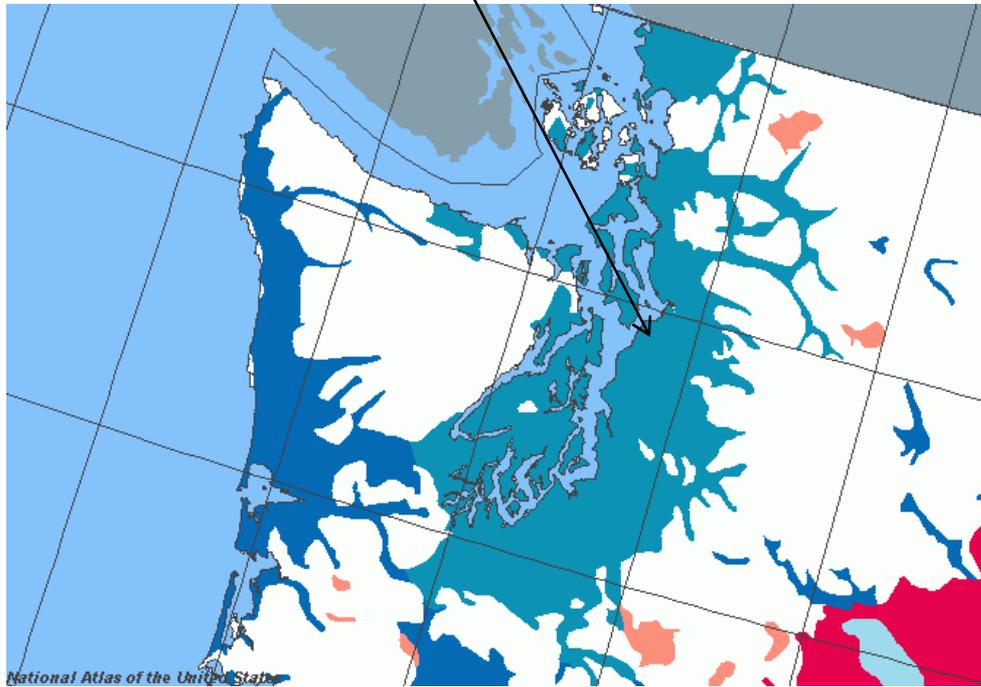


The average velocity of the water is the Darcy equation divided by the porosity of the sediment.



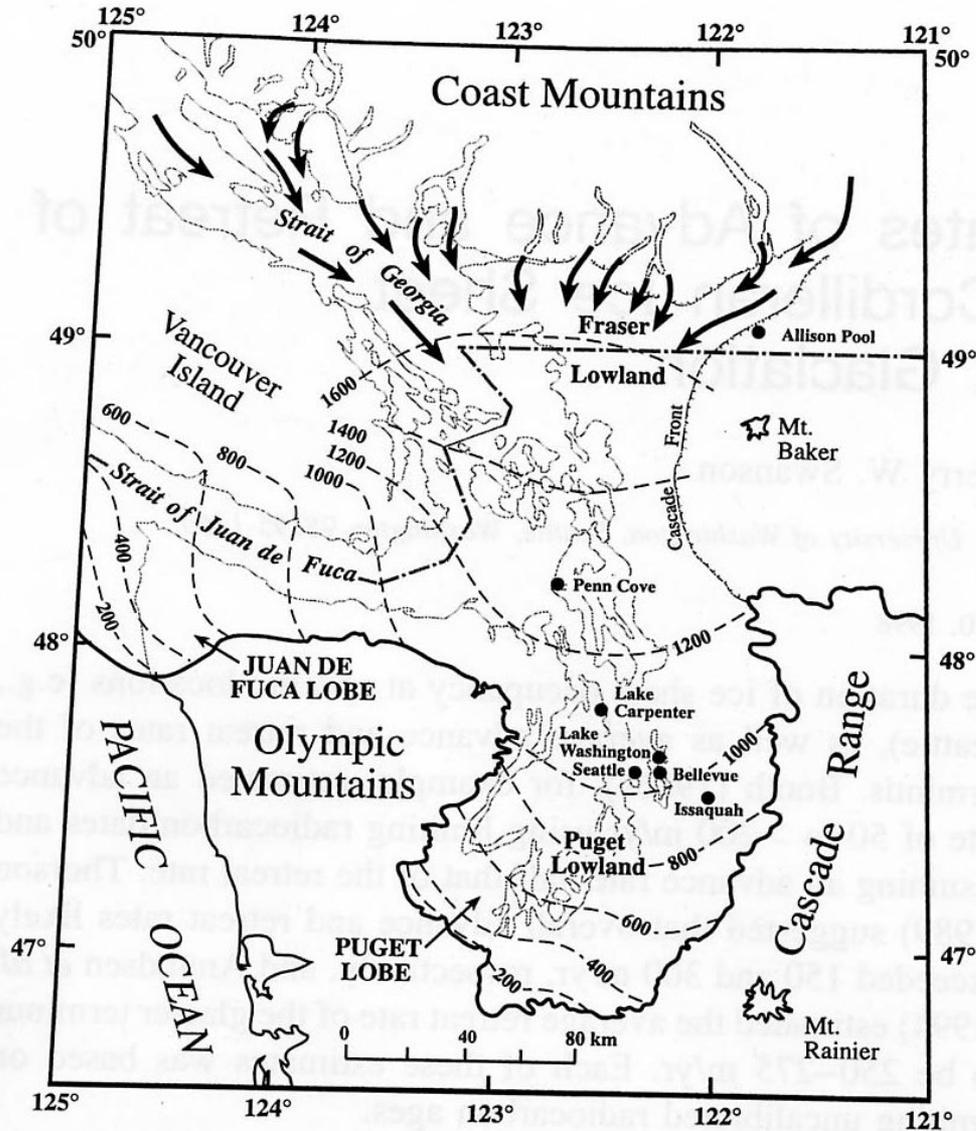
average pore water velocity  $v = -K/n(\Delta h/\Delta L)$

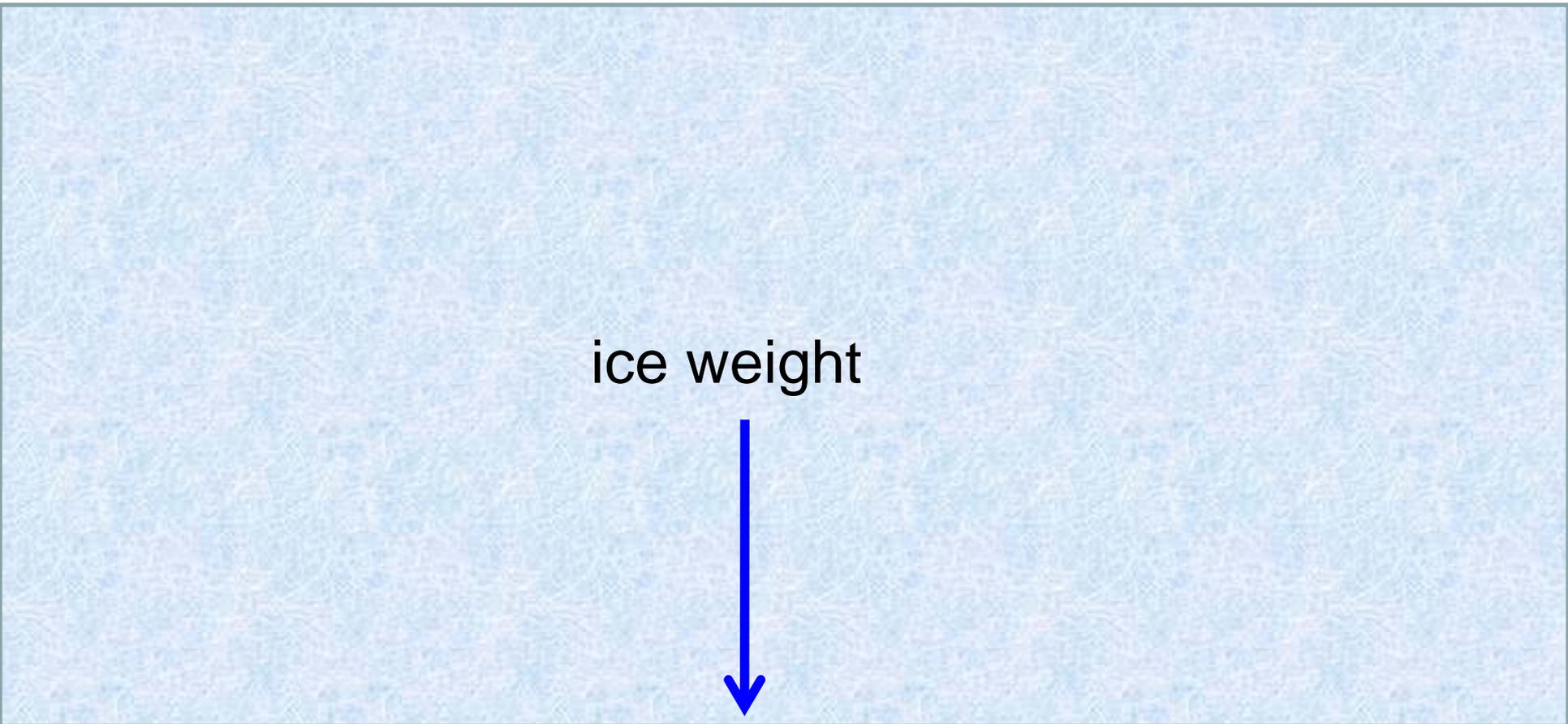
# Most aquifers in the Puget Lowlands are unconsolidated glacial deposits



[http://pubs.usgs.gov/ha/ha730/ch\\_h/H-text10.html](http://pubs.usgs.gov/ha/ha730/ch_h/H-text10.html)

# Glacial Deposits

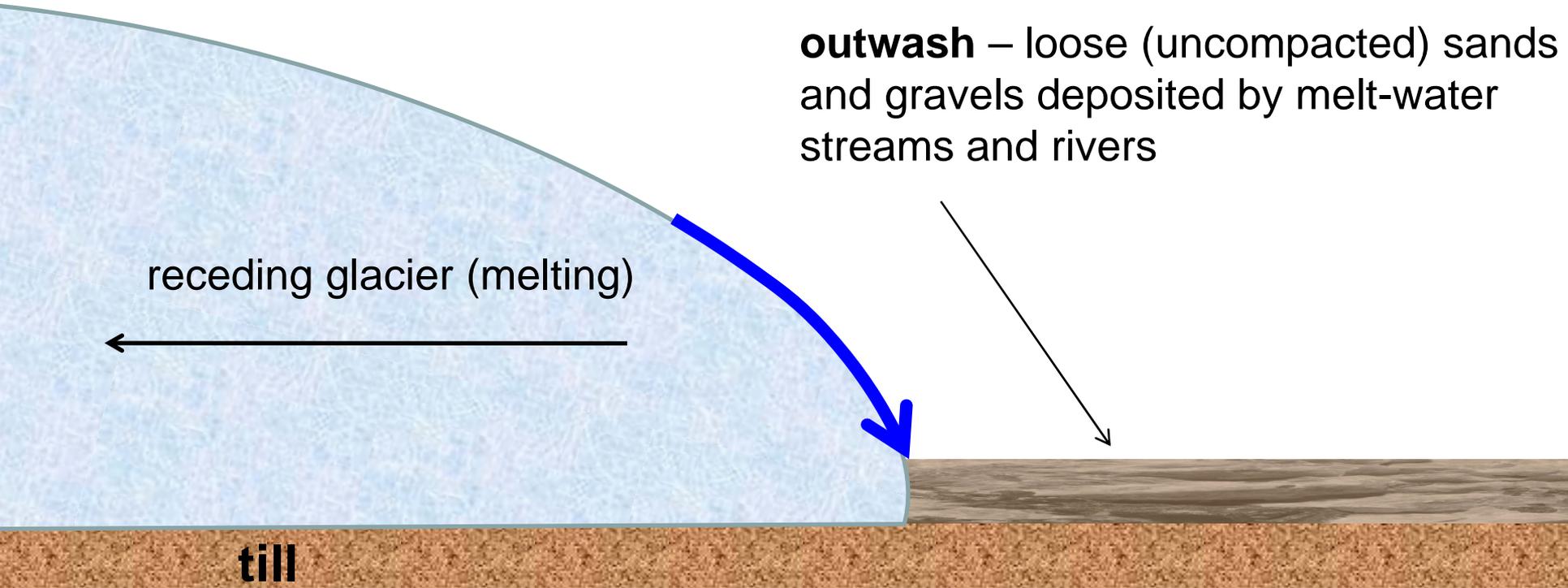




ice weight



**Till** – deposited beneath a glacier, it is a well compacted mixture of grain sizes  
(till does NOT make good aquifers)

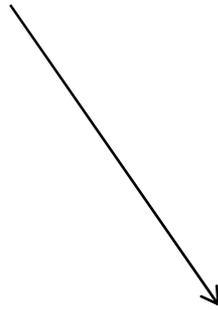


receding glacier (melting)

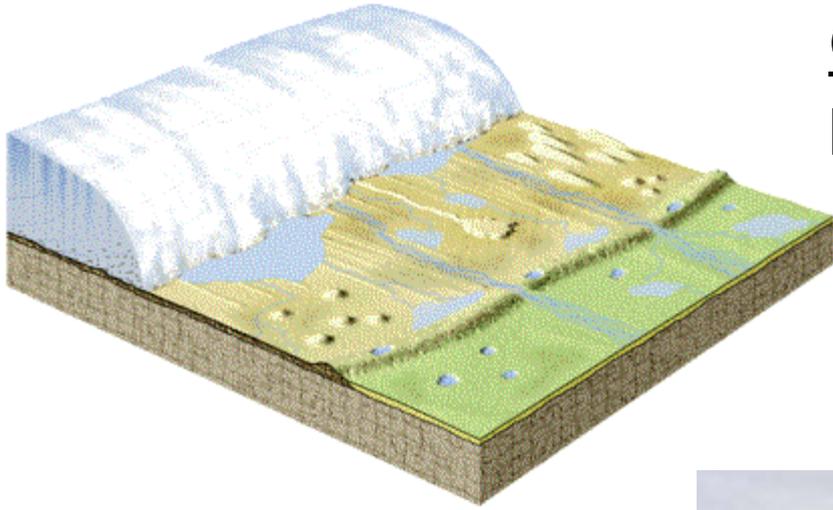


**till**

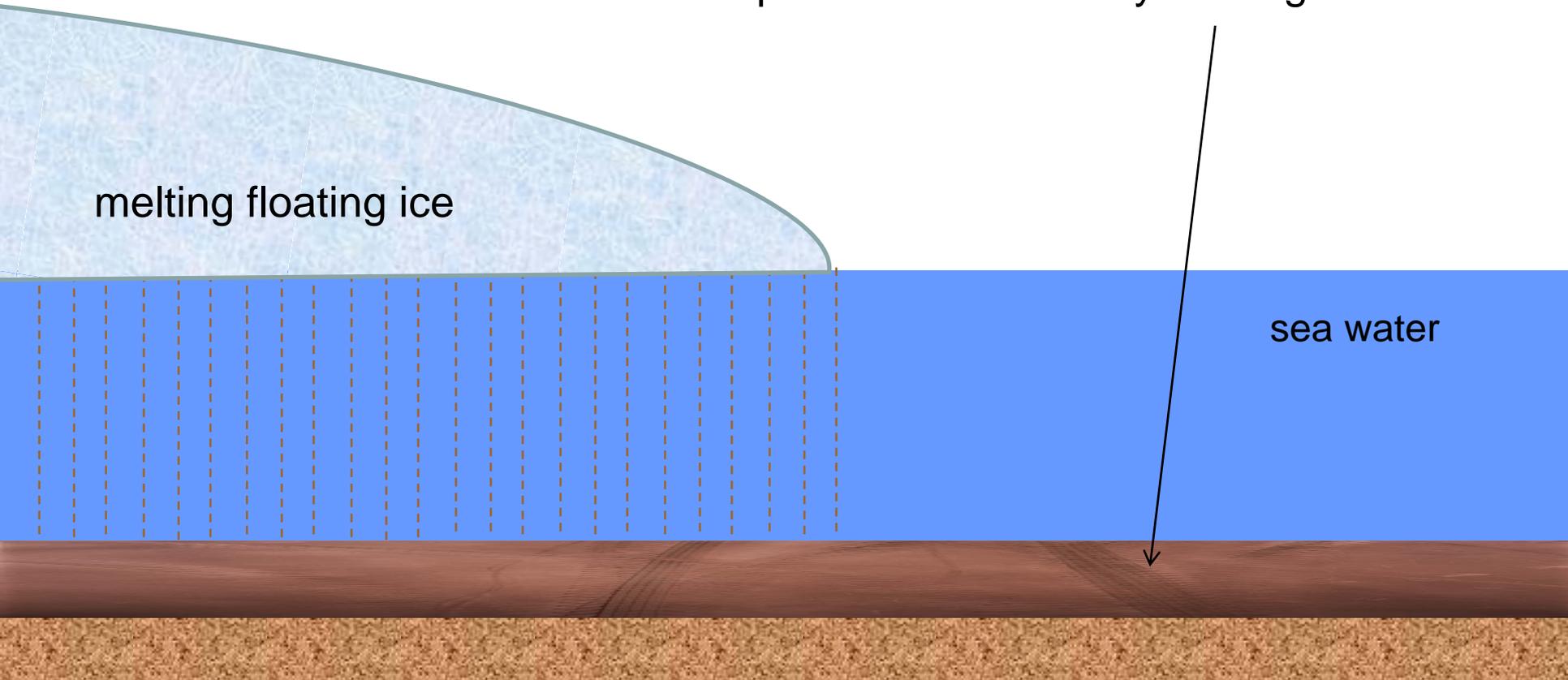
**outwash** – loose (uncompacted) sands and gravels deposited by melt-water streams and rivers



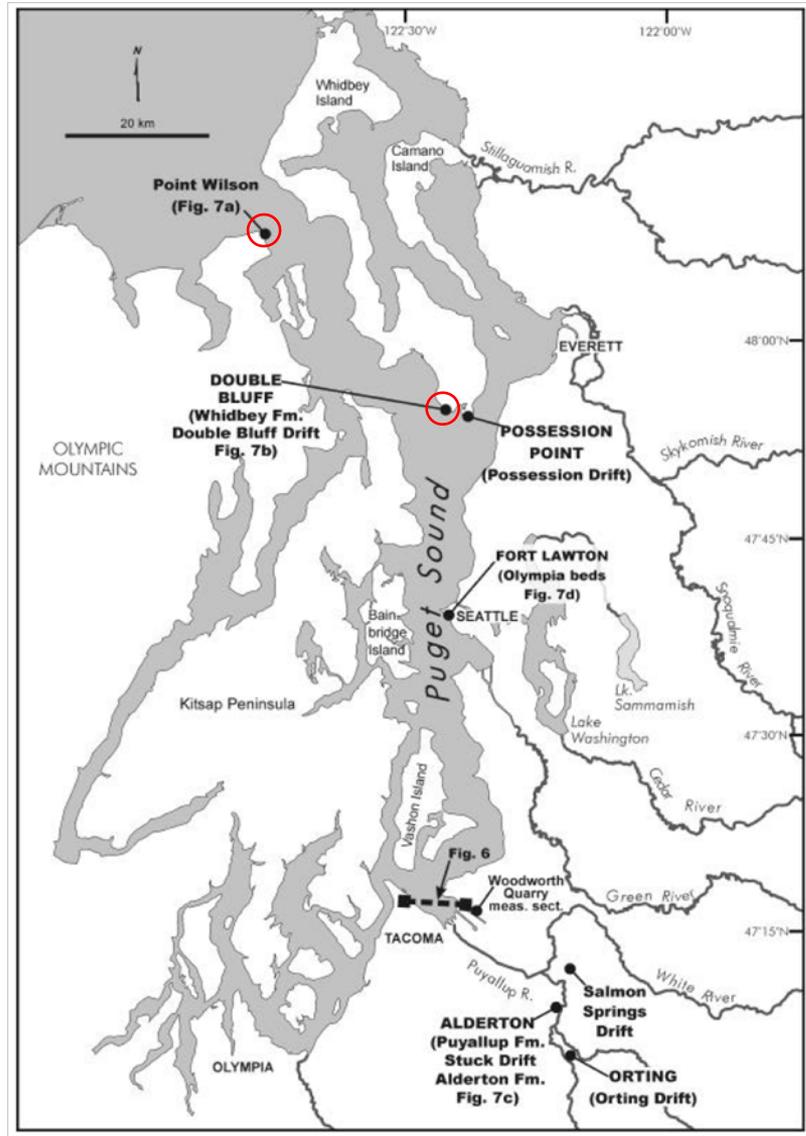
Glacial outwash makes the most productive aquifers



**glaciomarine drift** – clay rich sediment deposited in the sea by melting ice



glaciomarine drift does NOT make good aquifers

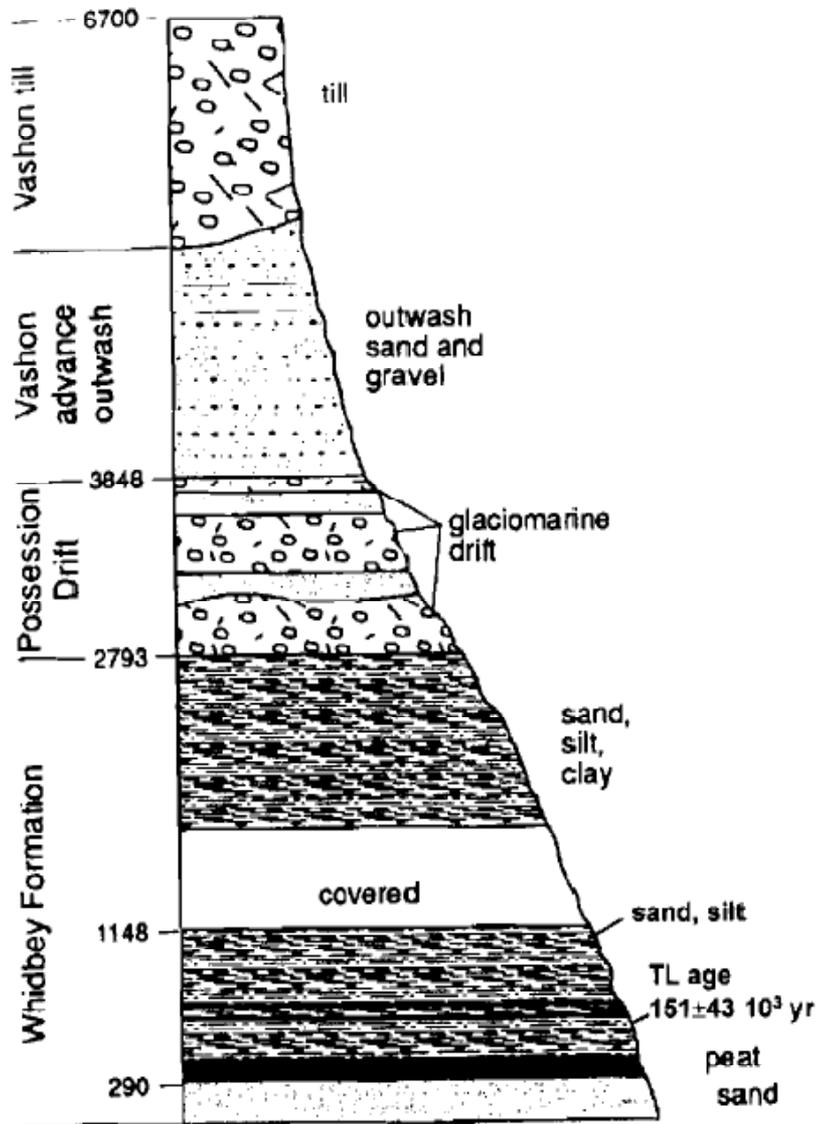


### THE CORDILLERAN ICE SHEET

by Derek B. Booth, Kathy Goetz Troost, John J. Clague, and Richard B. Waitt

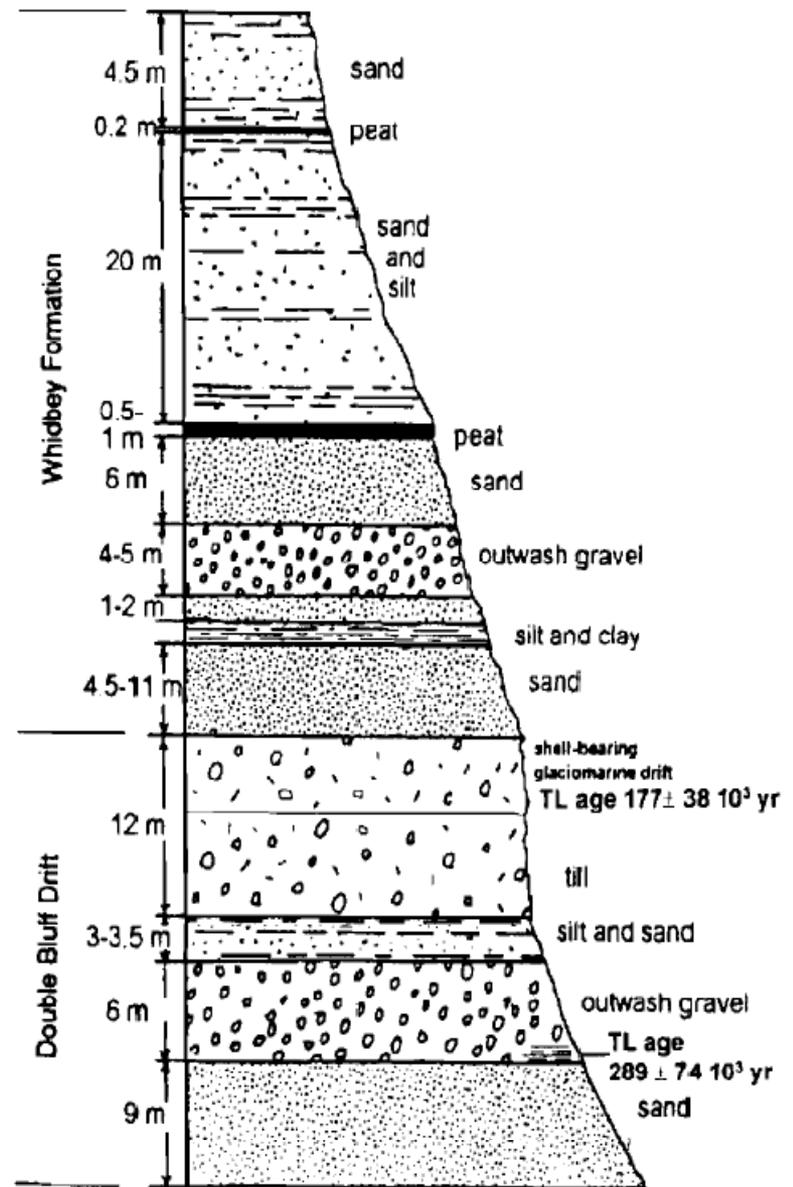
Preprint of Chapter 3 from *The Quaternary Period in the United States*  
 To be published by the International Union for Quaternary Research  
 August 2003

# Point Wilson



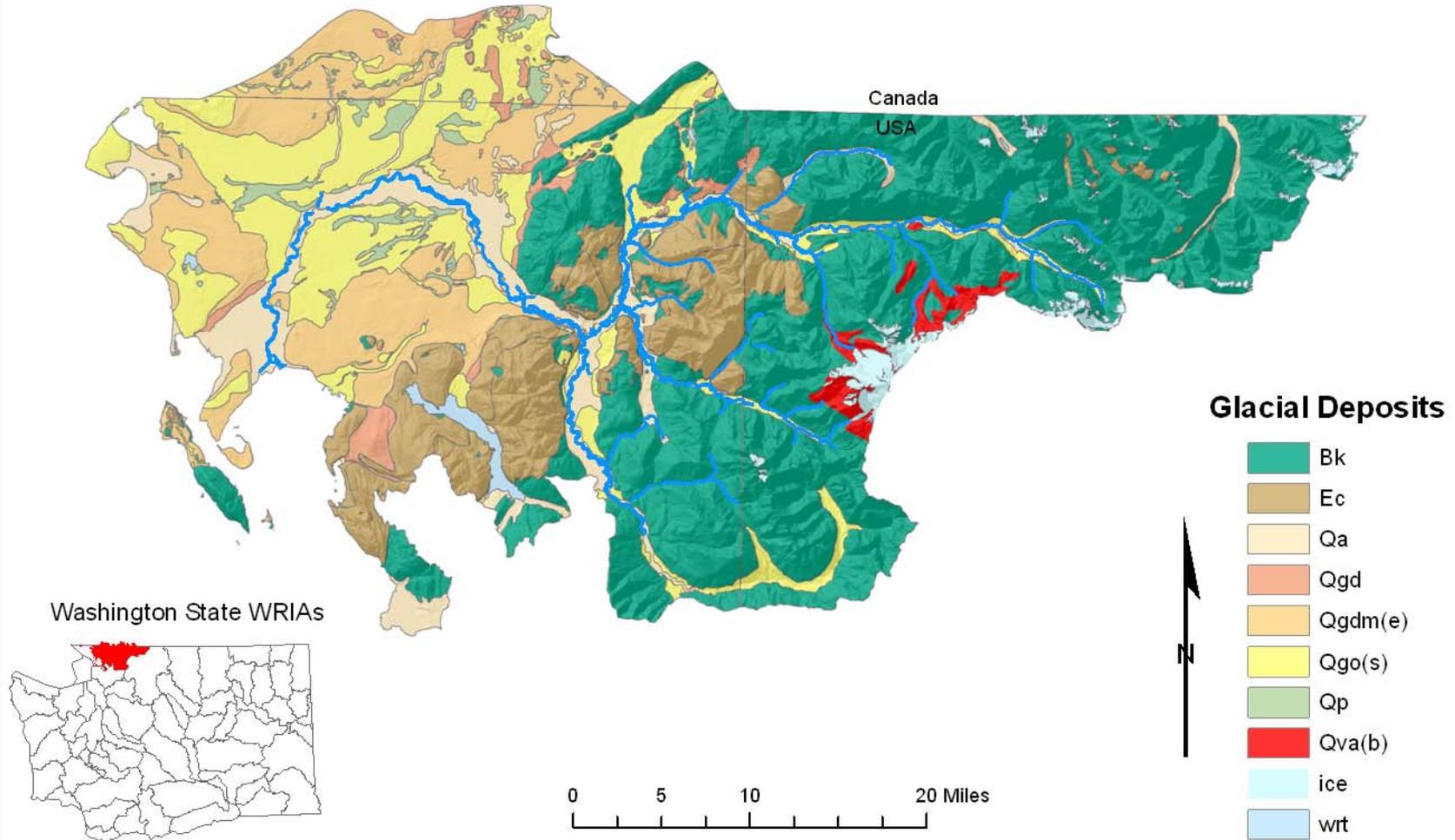
(a)

# Double Bluff



(b)

# Whatcom County Glacial Aquifers



# Whatcom County Glacial Deposits

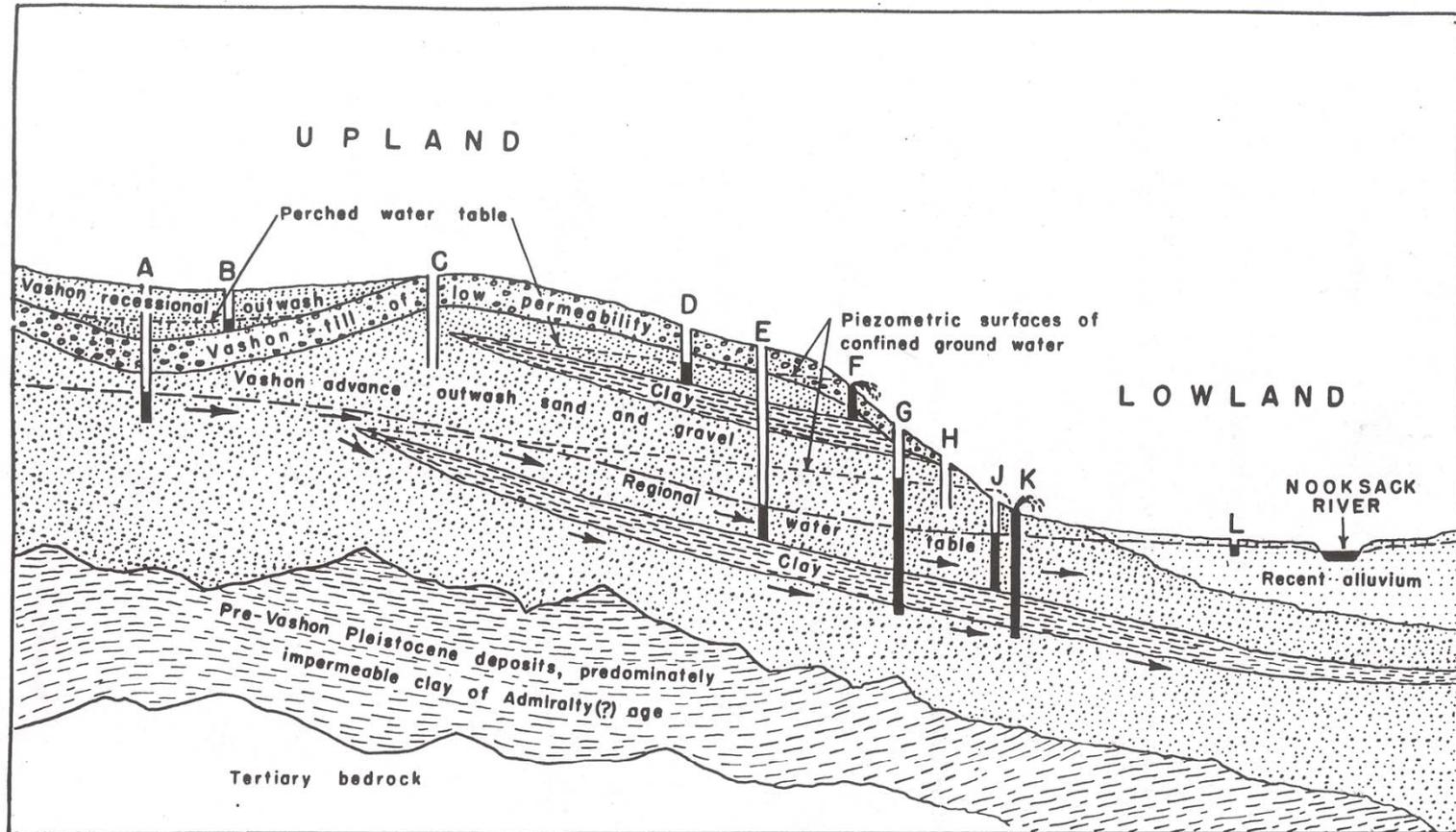
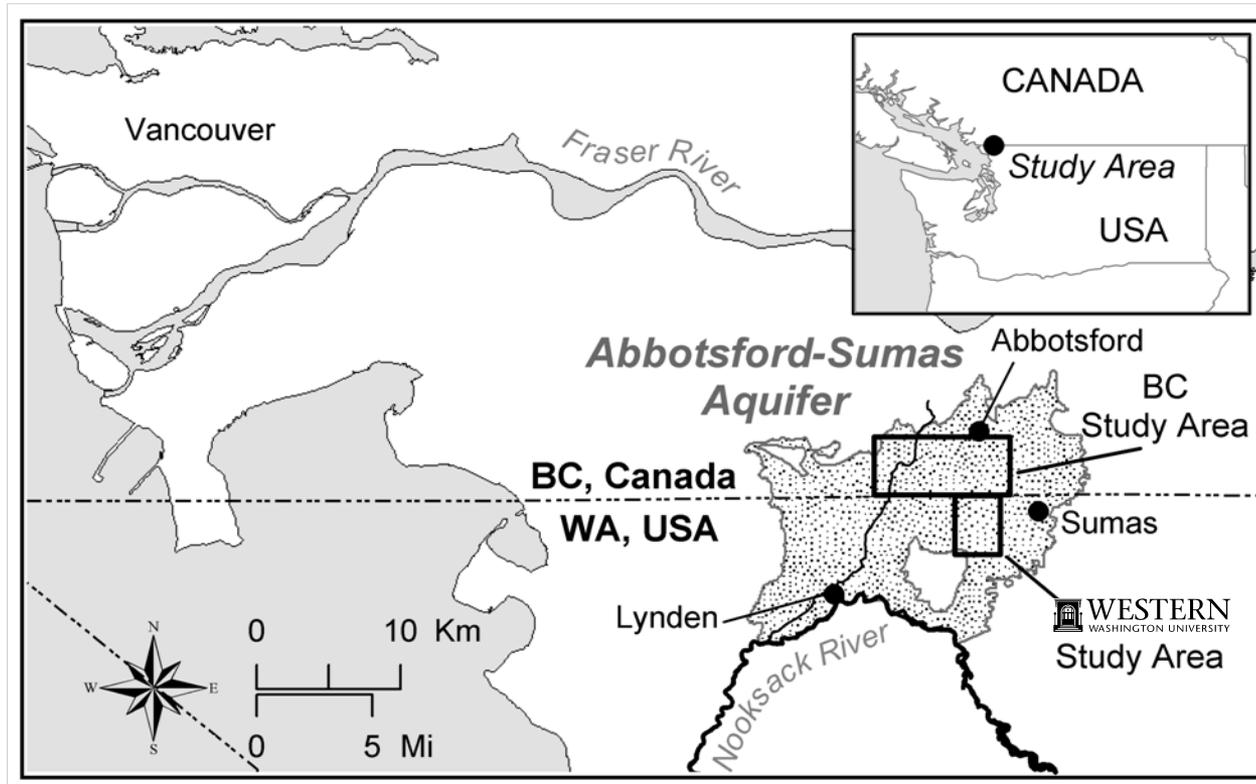


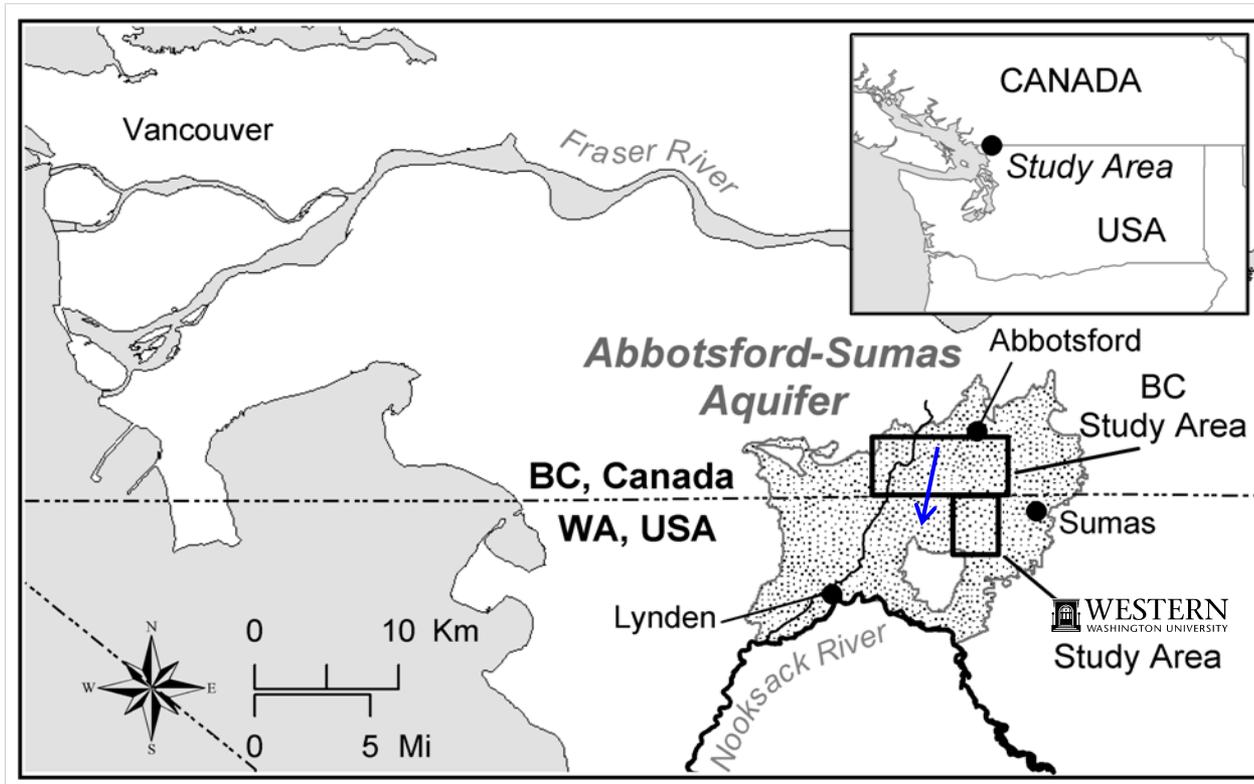
Figure 11. GENERAL CONDITIONS OF GROUND-WATER OCCURRENCE WITHIN THE LOWER NOOKSACK RIVER BASIN. Vertical scale exaggerated. Arrows indicate direction of ground-water movement.

# Abbotsford-Sumas Aquifer

The aquifer covers approximately 200 km<sup>2</sup> and serves as a water supply for approximately 110,000 people in BC and WA.

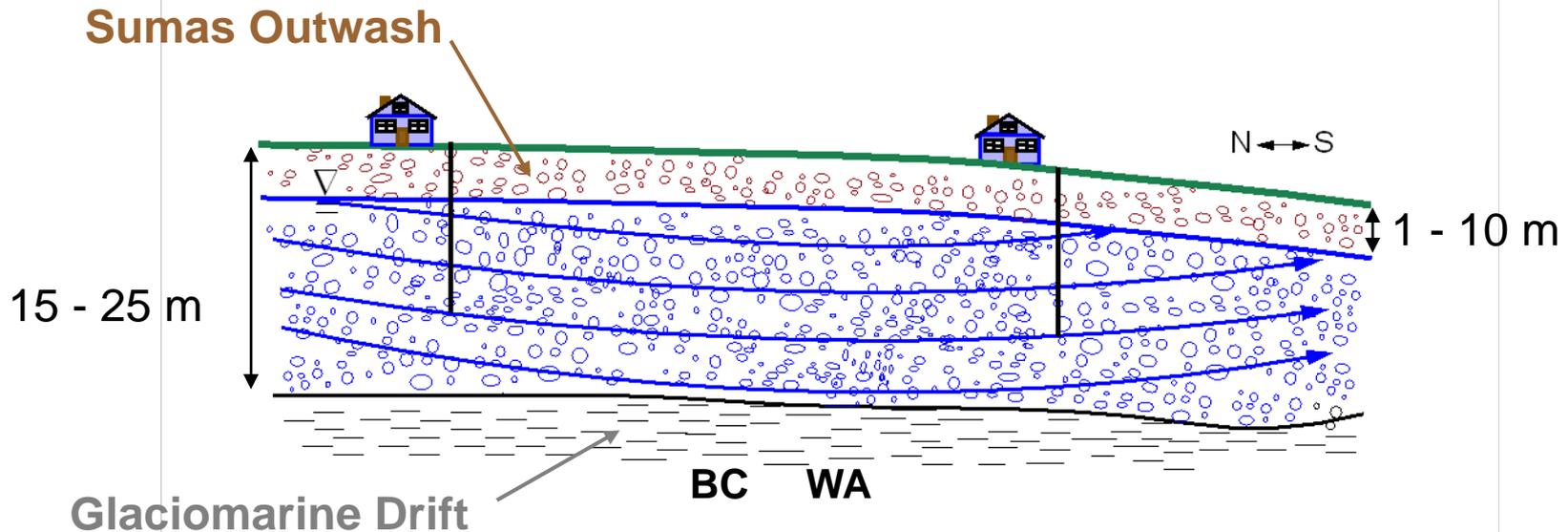


## Ground Water Flow is from North-to-South

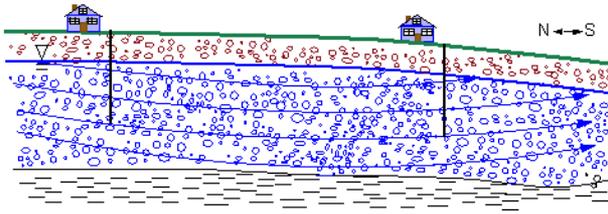


# Abbotsford-Sumas Aquifer

The aquifer is unconfined and comprised of glacial outwash sands and gravels (Sumas Outwash) deposited about 10,000 years ago.



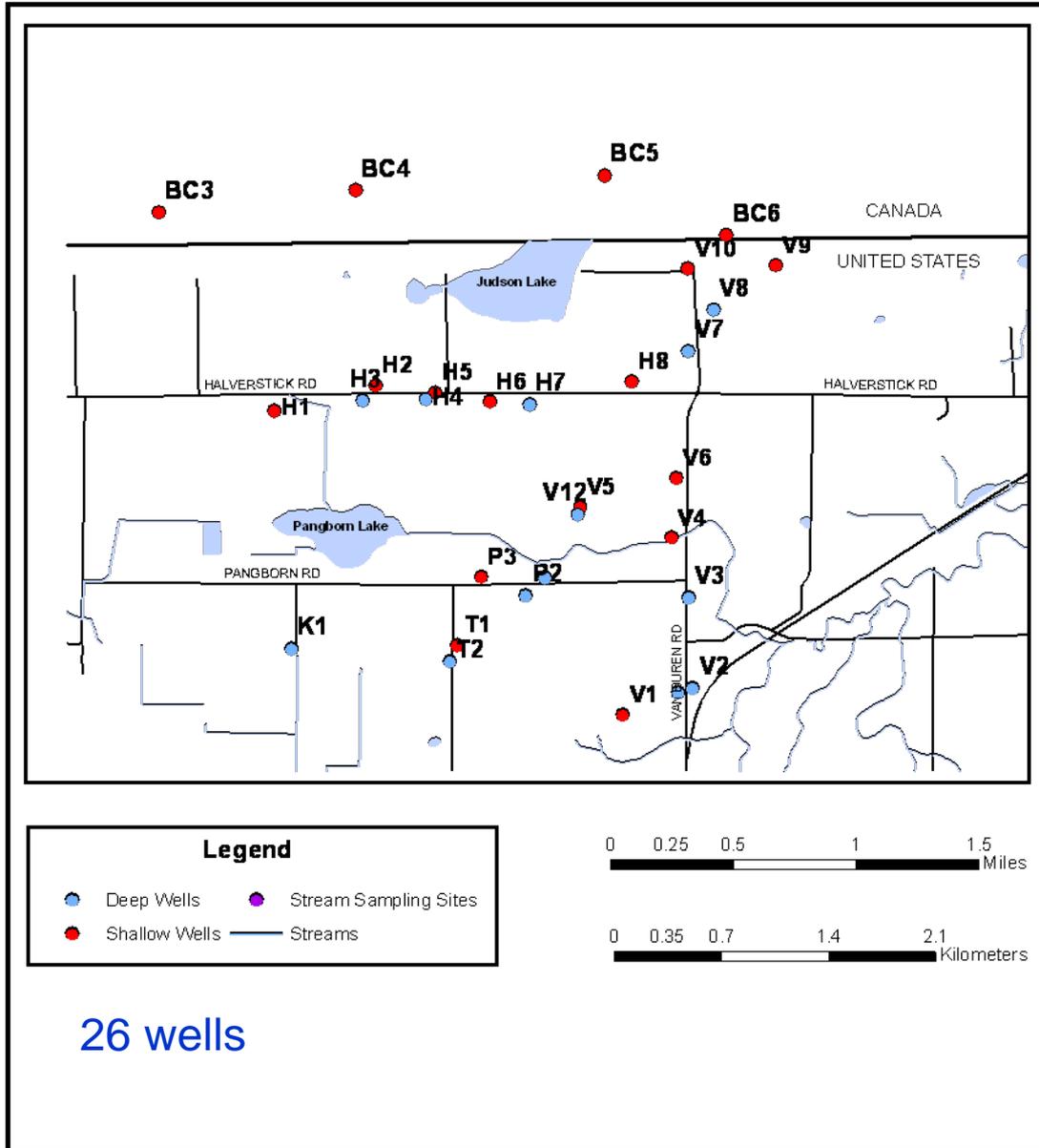
# Sumas Outwash



gravel mine in Whatcom County



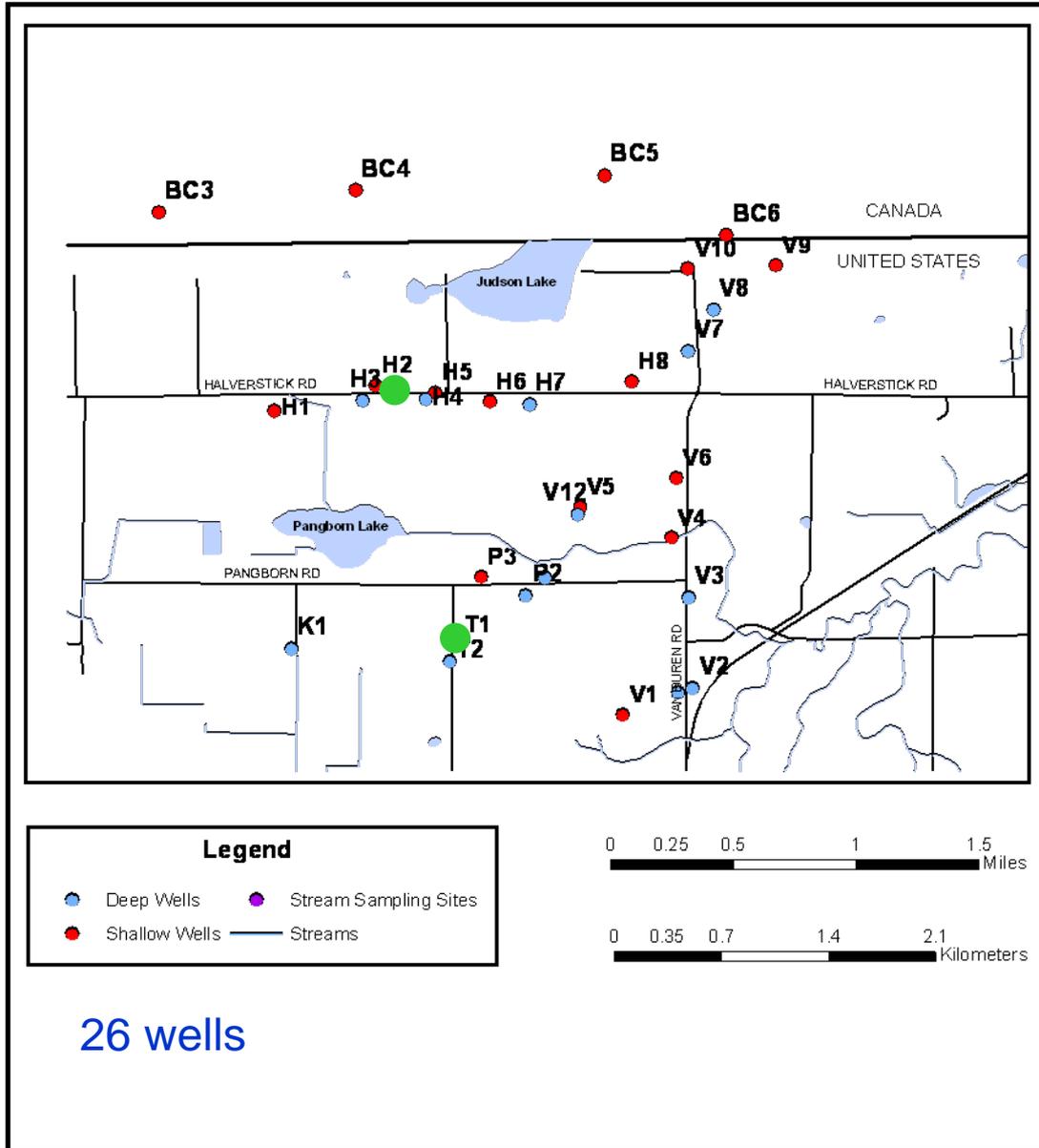
# Well Sampling Sites



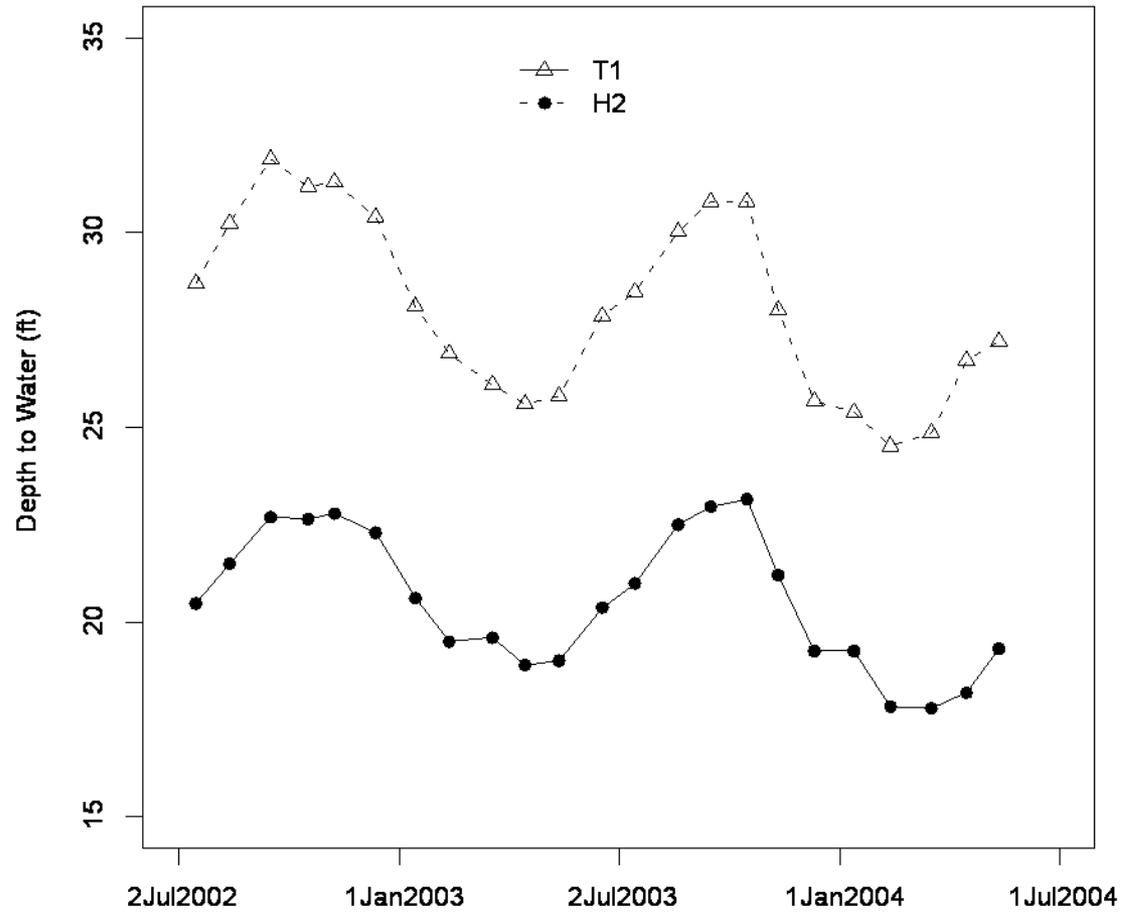
# Water-level measurements



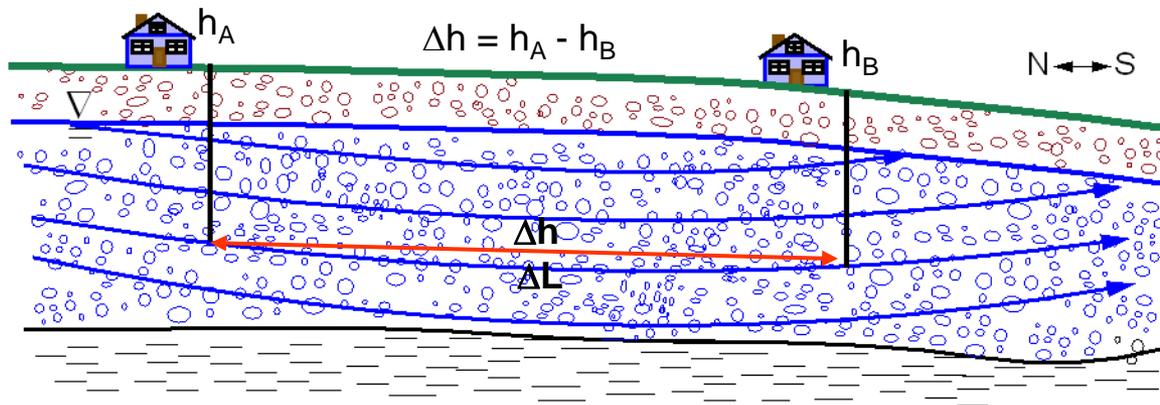
# Well Sampling Sites



# Water Table Hydrographs



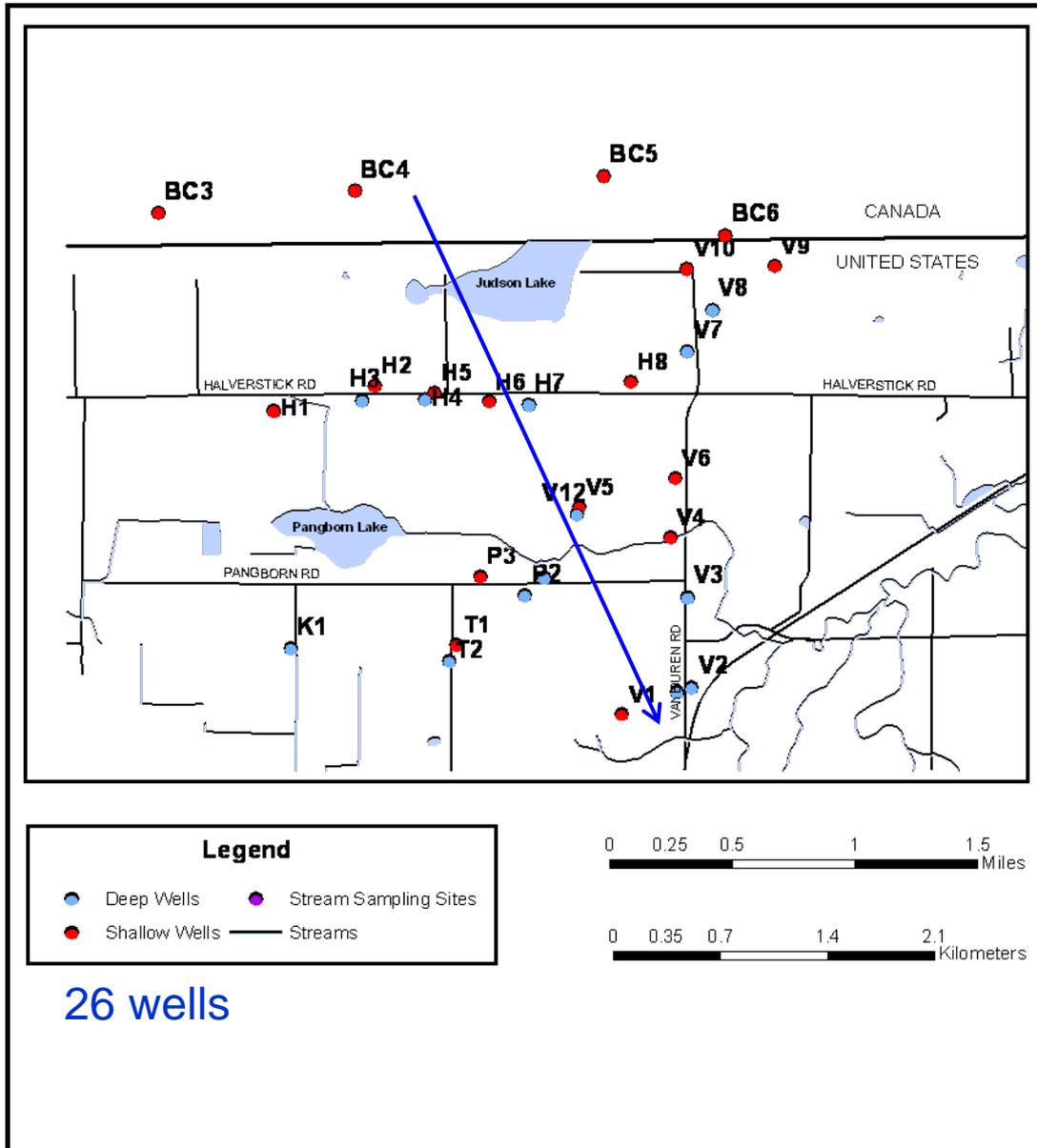
hydraulic gradient =  $\Delta h / \Delta L$



# Well Sampling Sites

Flow is in the southeast direction (155°) and is drive by a gradient of

$$\frac{\Delta h}{\Delta L} = -0.0055$$



The average **hydraulic gradient** in the Abbotsford-Sumas aquifer is

$$\Delta h/\Delta L = -0.0055$$

The average **hydraulic conductivity** of the glacial sediments is

$$K = 545 \text{ ft/day} \quad \text{or} \quad K = 0.187 \text{ cm/sec}$$

The average **porosity** of the glacial sediments is

$$n = 0.30$$

The average pore-water velocity can be determined using

$$\text{velocity} = v = -K/n (\Delta h/\Delta L)$$

Using the aquifer parameters in the equation above yields

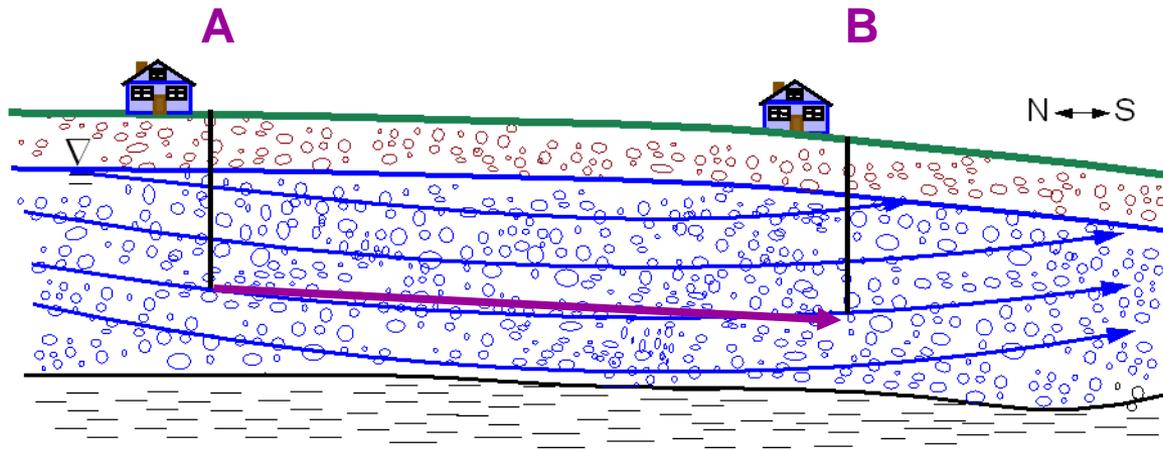
$$\text{velocity} = v = -545/0.30(-0.0055) = 10 \text{ ft/day}$$

which is very fast for groundwater

The average groundwater velocity is 10 ft/day

How long would it take water to travel from well A to well B

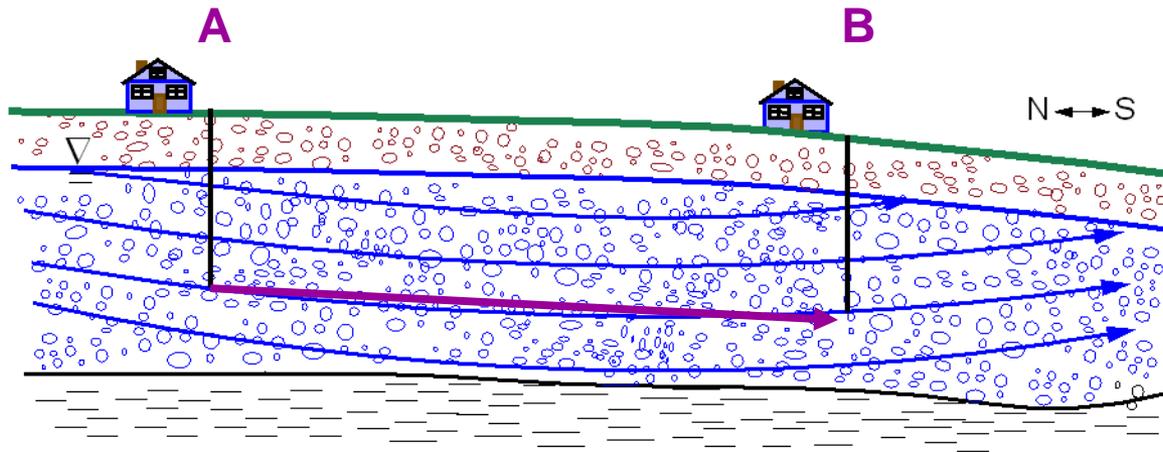
If the distance from A to B is 1000 ft



Time = distance / velocity

Time = 1000 ft / 10 ft/day

Time = 100 days

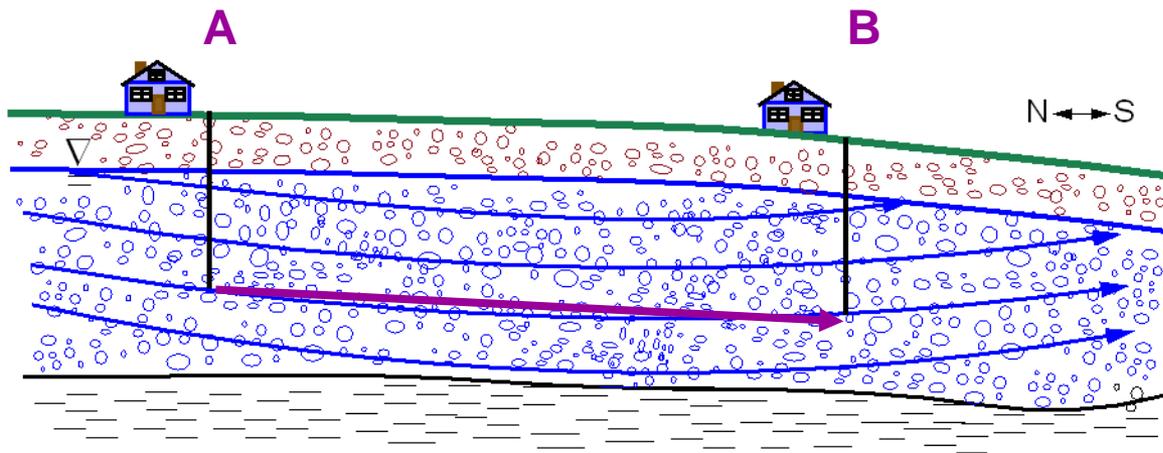


If the aquifer were a fine sand with a hydraulic conductivity of 5.45 ft/day.....then the

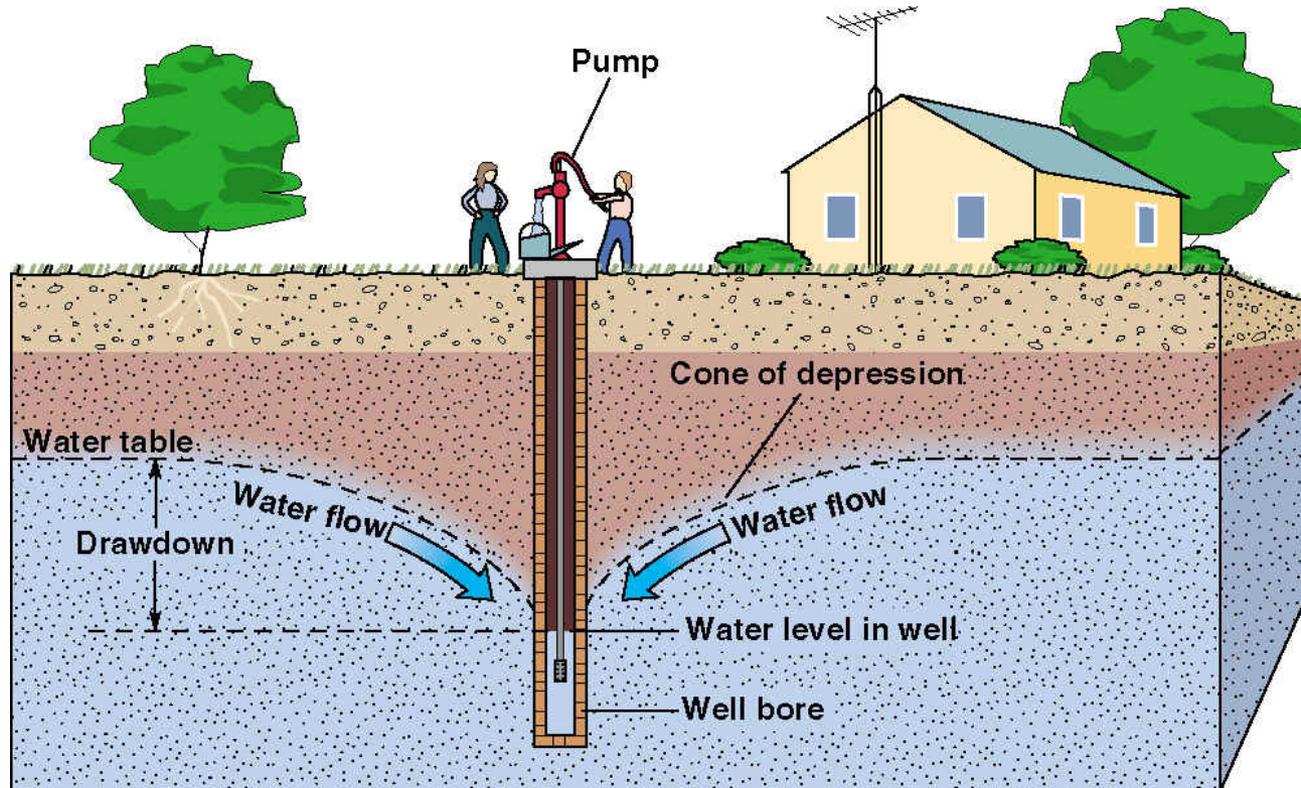
Velocity = 0.10 ft/day

Time = 1000 ft / 0.10 ft/day

Time = 10,000 days or 27 years!!

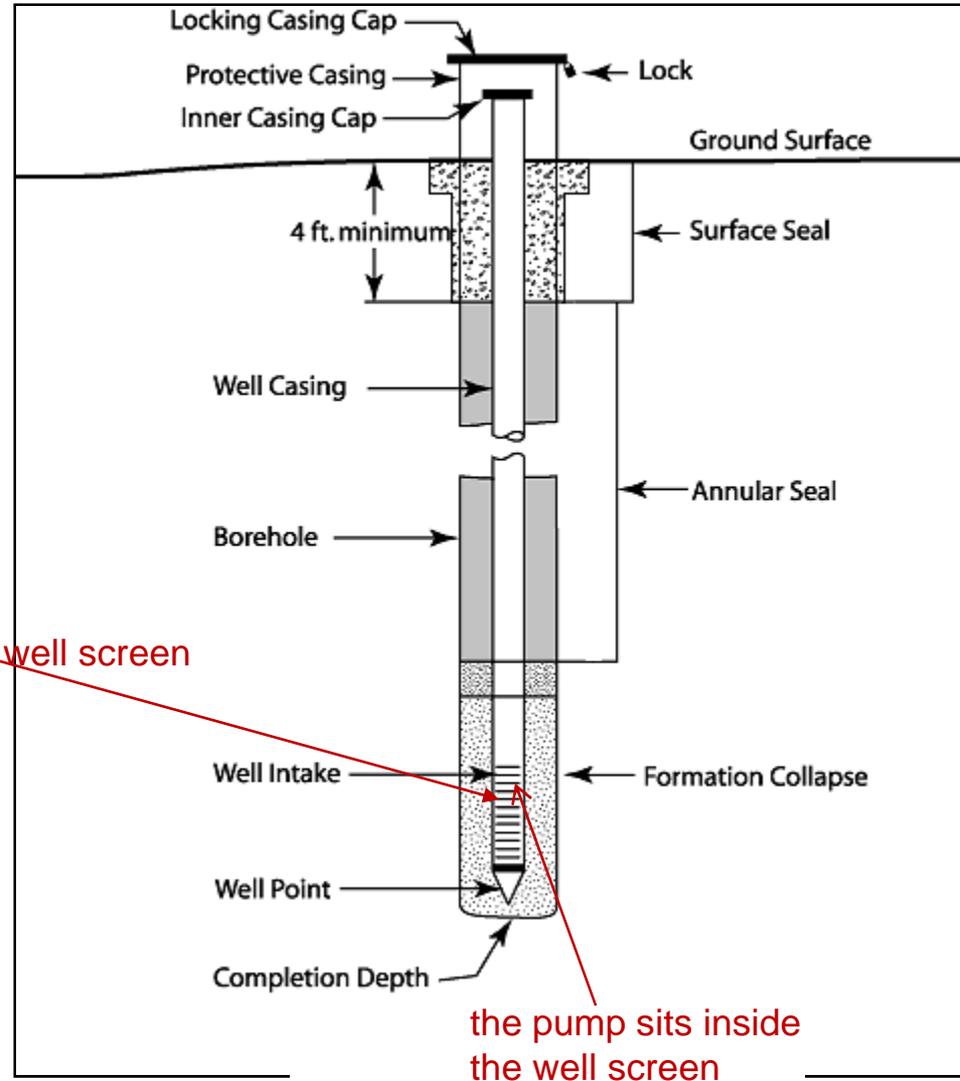
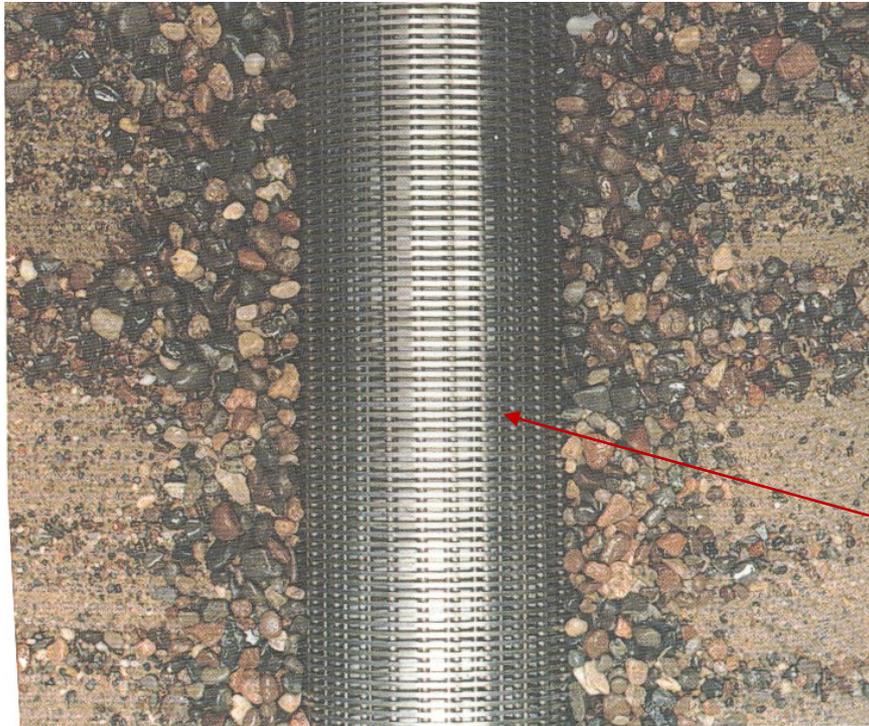


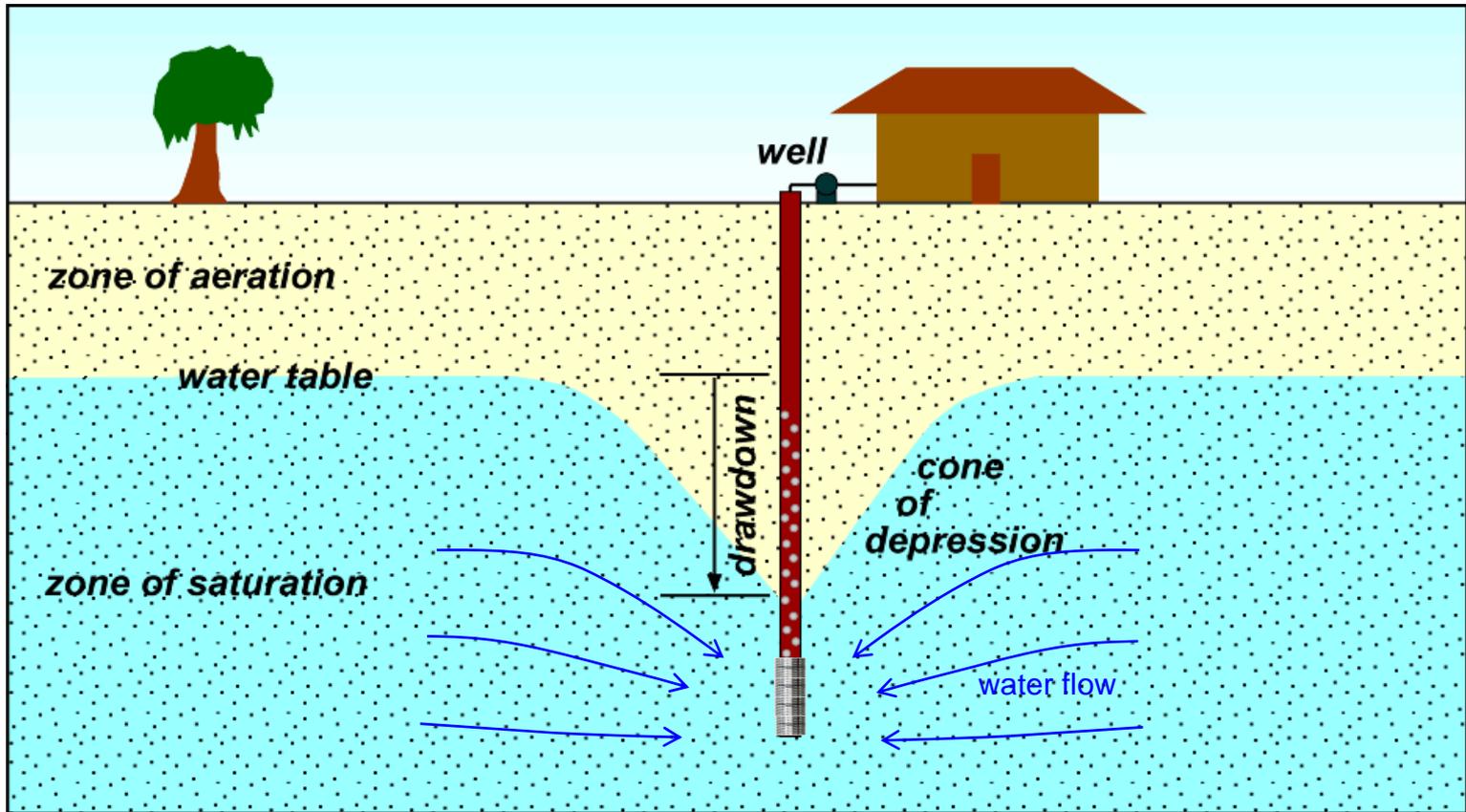
# Pumping wells create a cone-of-depression in the water table

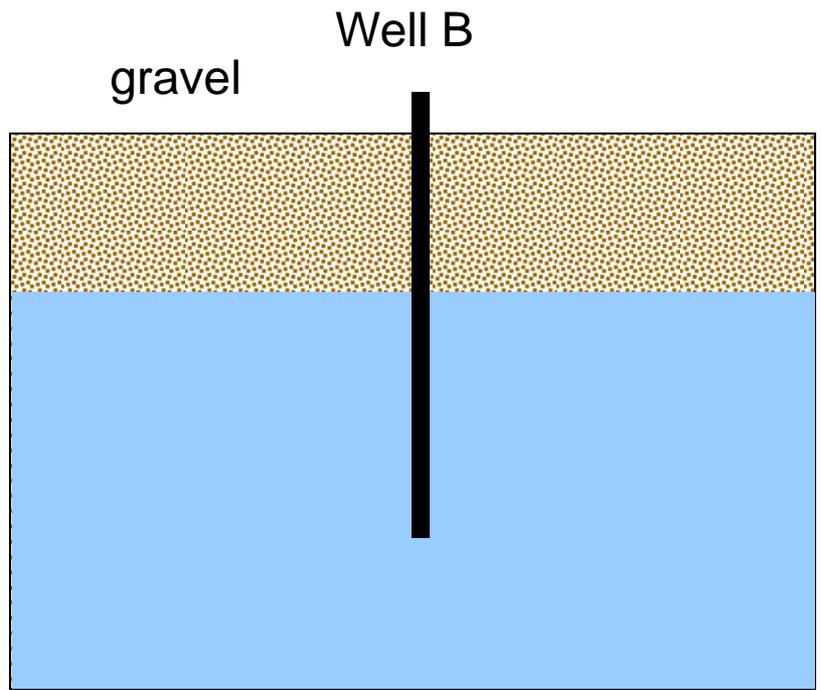
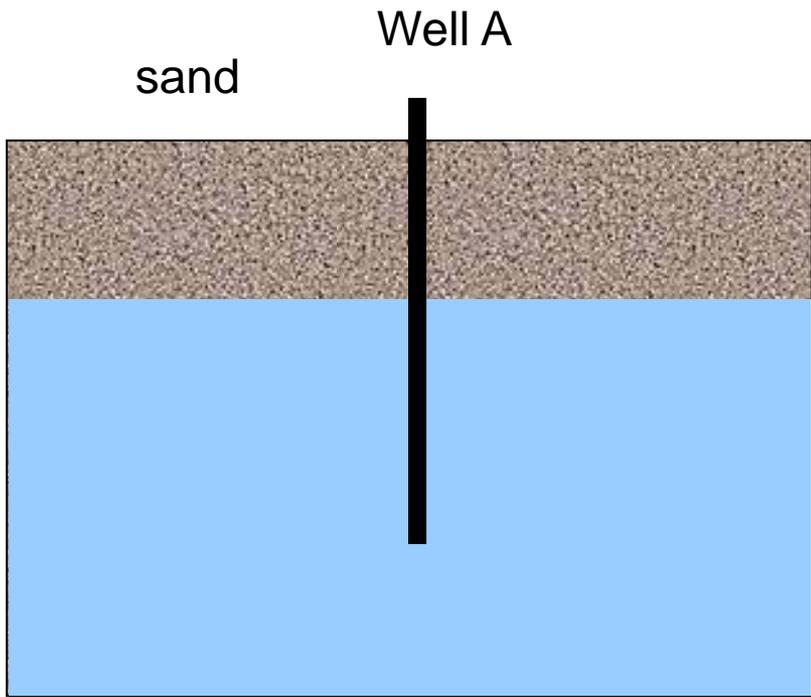


© 2001 Brooks/Cole - Thomson Learning

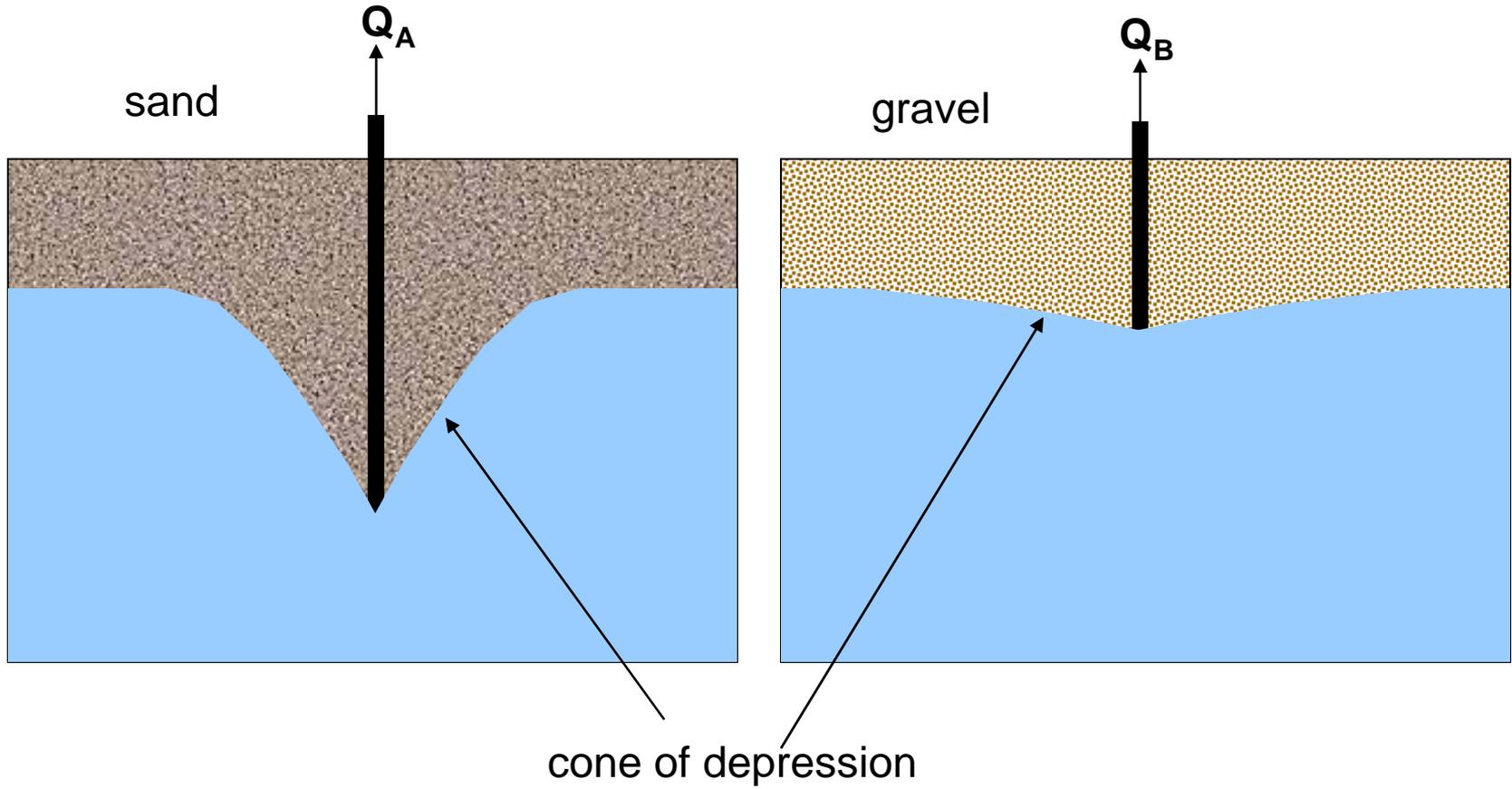
[http://www.uwsp.edu/geo/faculty/ozsvath/images/cone\\_of\\_depression.htm](http://www.uwsp.edu/geo/faculty/ozsvath/images/cone_of_depression.htm)

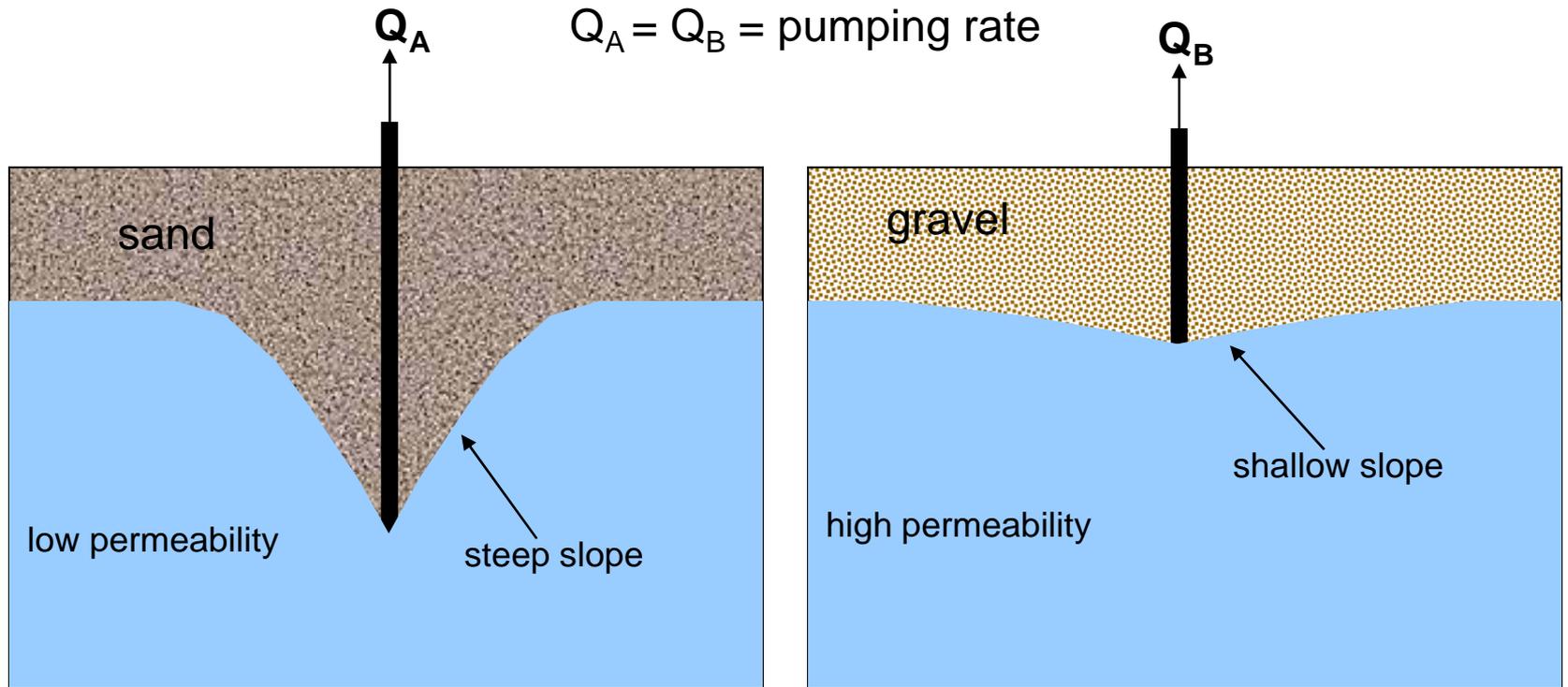




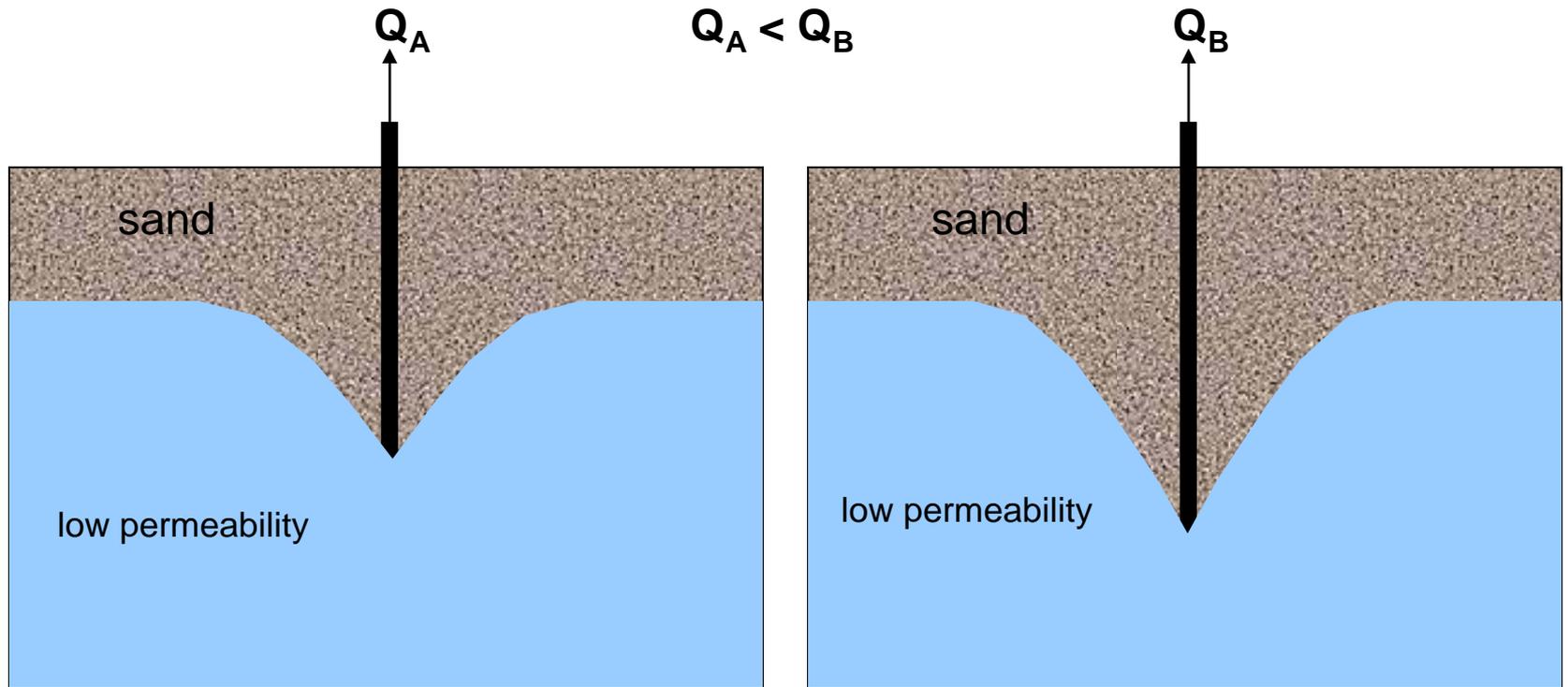


$$Q_A = Q_B = \text{pumping rate}$$

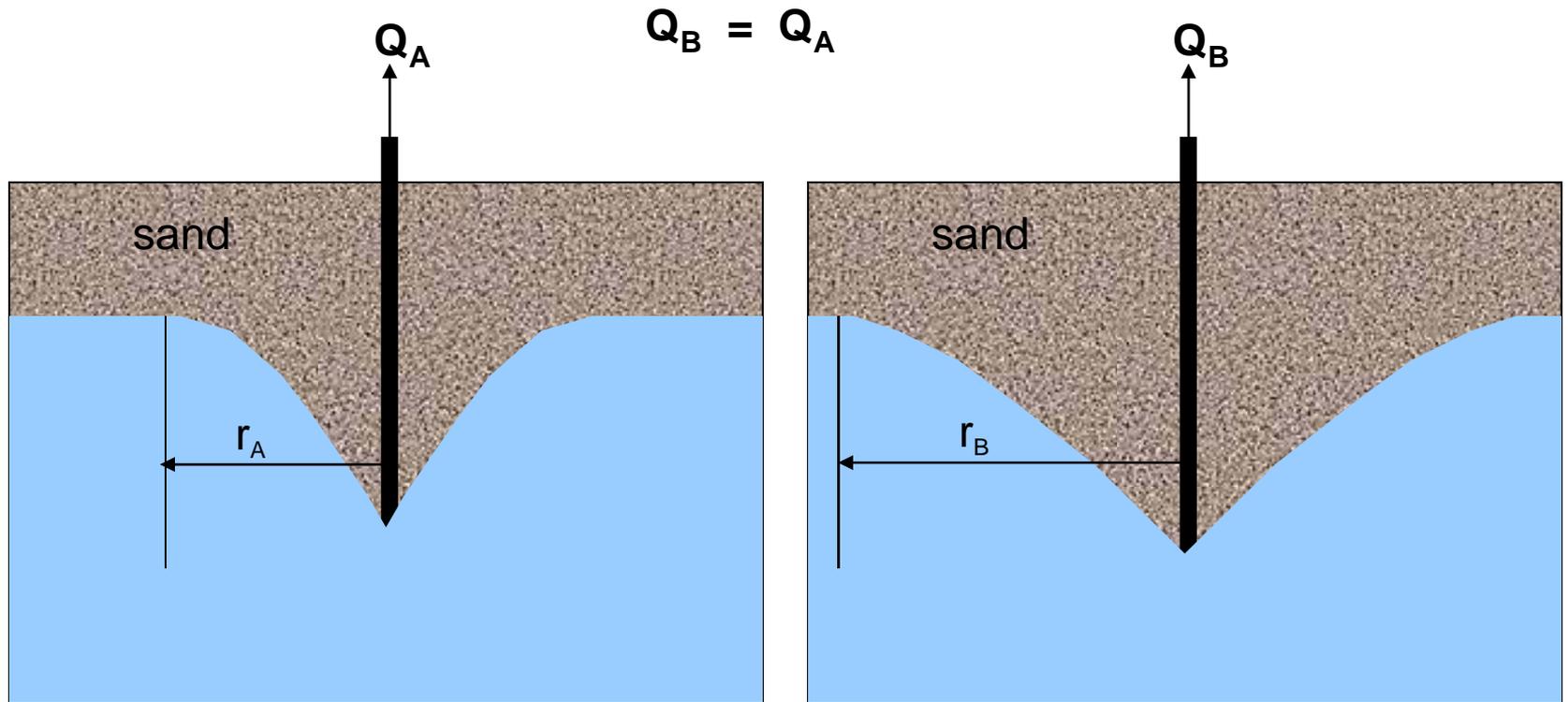




The slope of the cone of depression is determined by the permeability



The depth of the cone of depression is determined by the pumping rate



The radius of the cone of depression is determined by the pumping duration

$$r_A < r_B$$

Geotechnical, Rock and Water Resources Library - Grow Resource - Water Table Drawdown - Windows Internet Explorer

http://www.grow.arizona.edu/Grow-GrowResources.php?ResourceId=168

File Edit View Favorites Tools Help

Geotechnical, Rock and Water Resources Library - Gr...

greater is the drawdown. when the drawdown causes the water table to come down to the level of the intake screen of the pump the pump cannot pump any more water and must be turned off till the water table raises or the pump needs to be placed at a lower level. Sometimes this will require that the well be drilled deeper.

**To conduct the experiment:**

1. Select soil type
2. Turn on the pump

### Water Table Drawdown Experiment

Gravel

Sand

Sandy-clay

Static Water Table

Aquifer

Well Casing

Power Switch

Pump

Screen intake

Motor

Reset

About

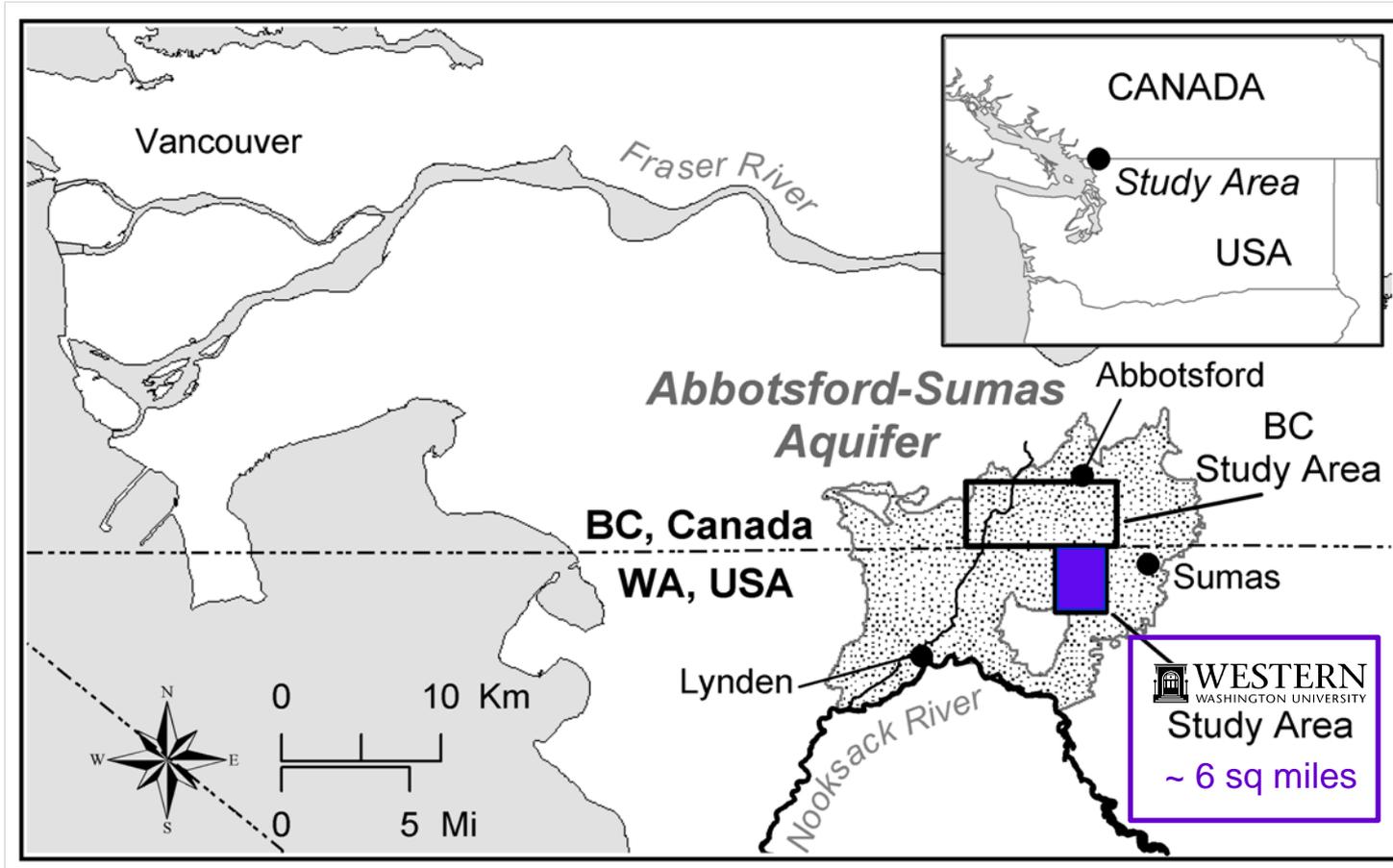
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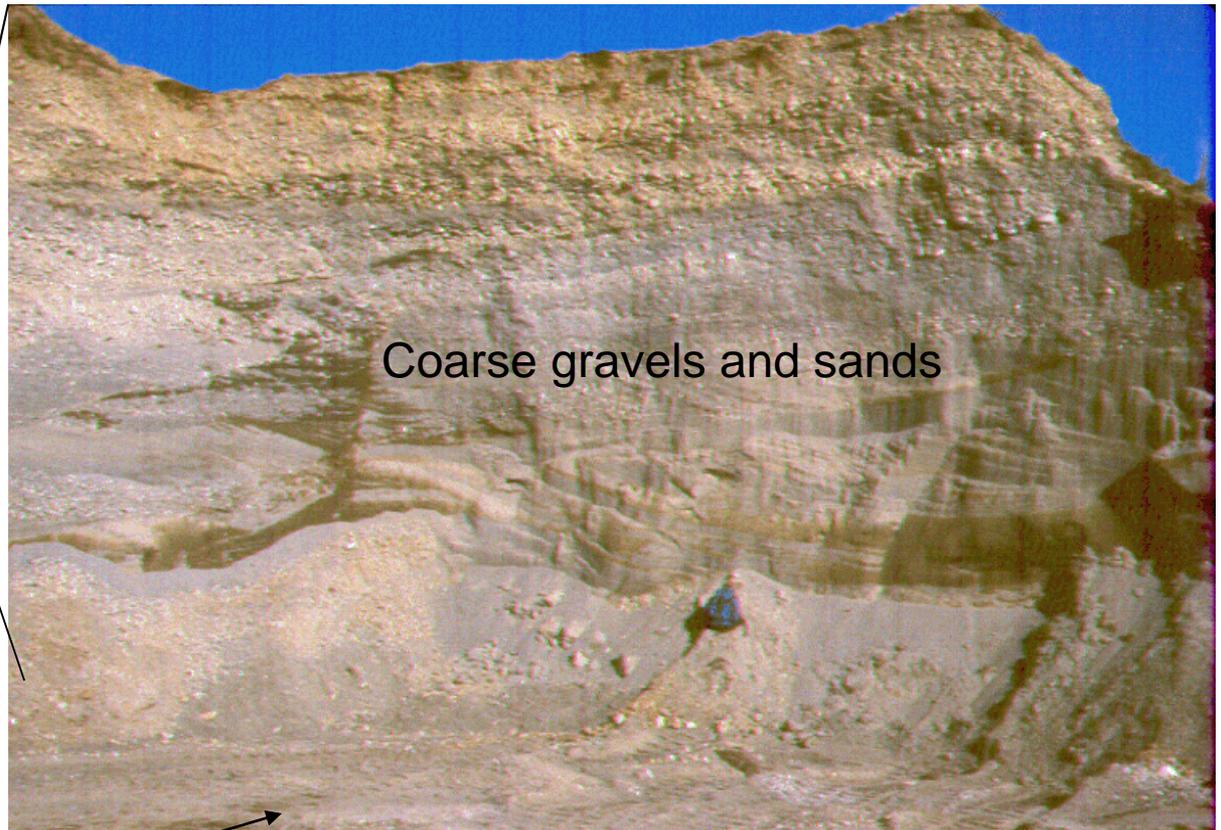
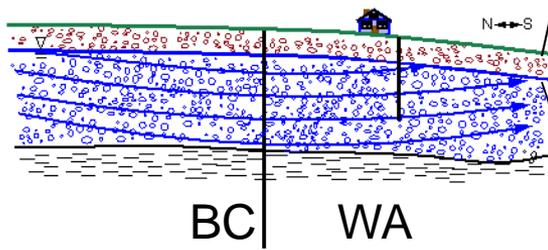
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Internet 125%

# Study Area

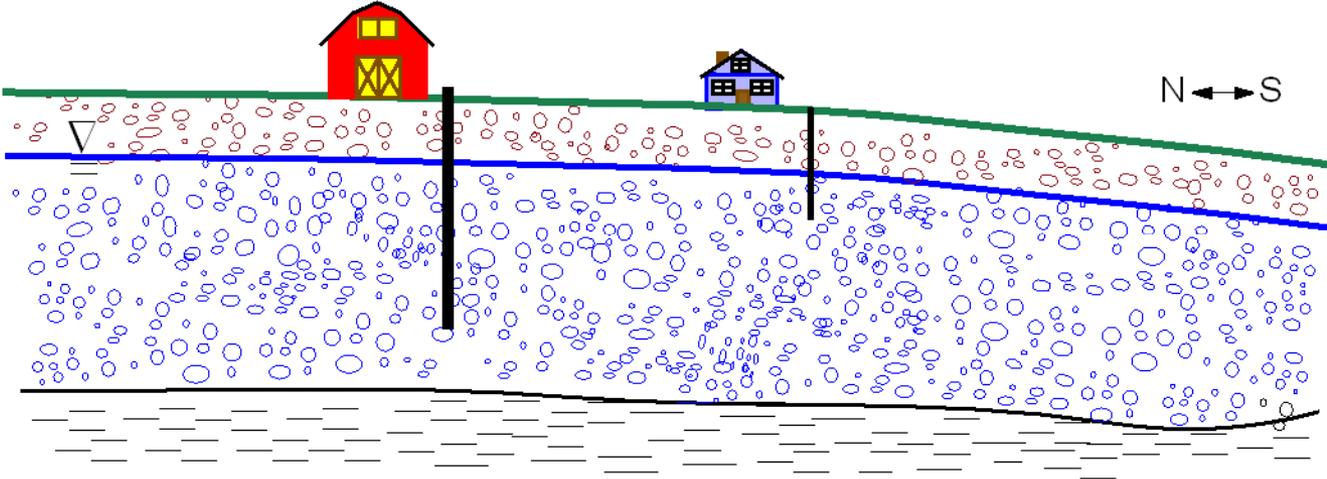




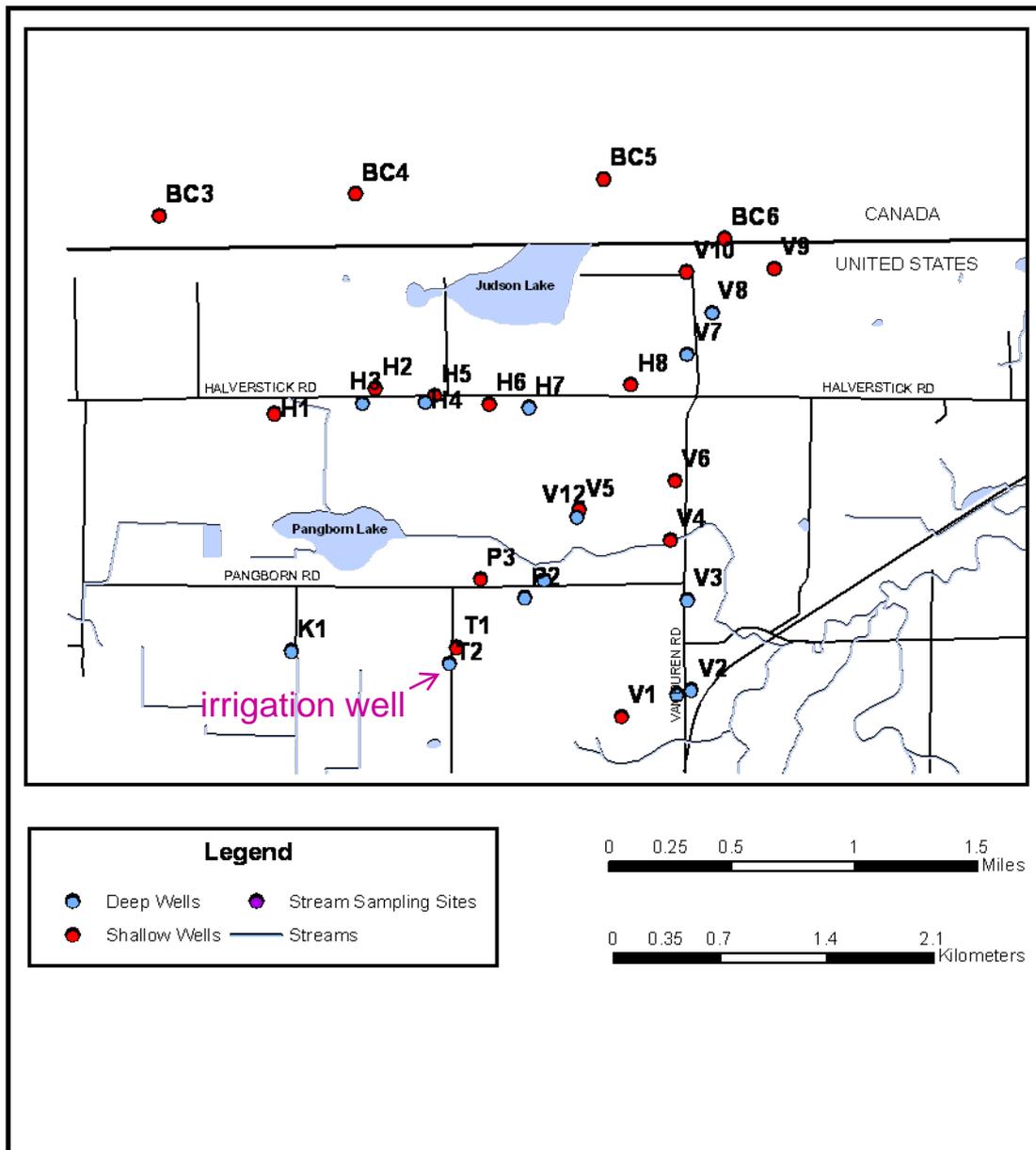
Coarse gravels and sands

Water table is just below the ground surface

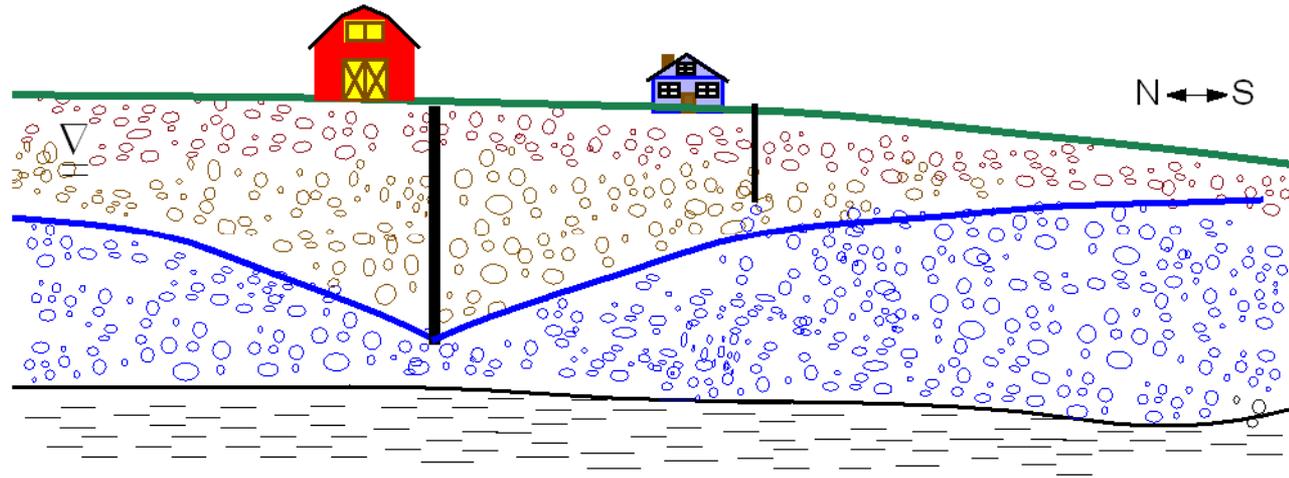
# Pumping wells can influence neighboring wells



# Irrigation Well

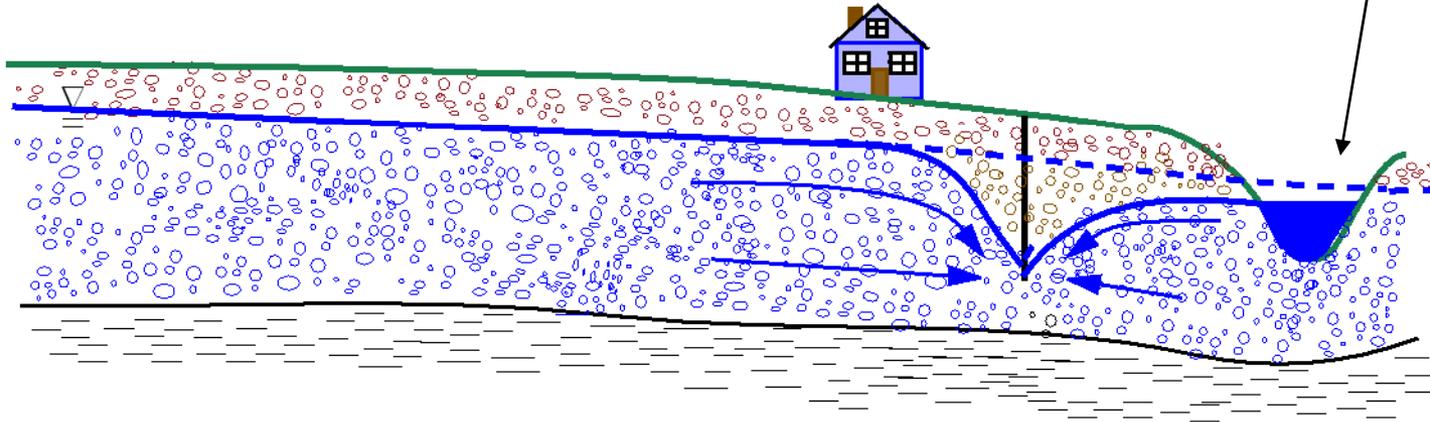


Over pumping of irrigation well lowers the water table below the domestic well (water rights)

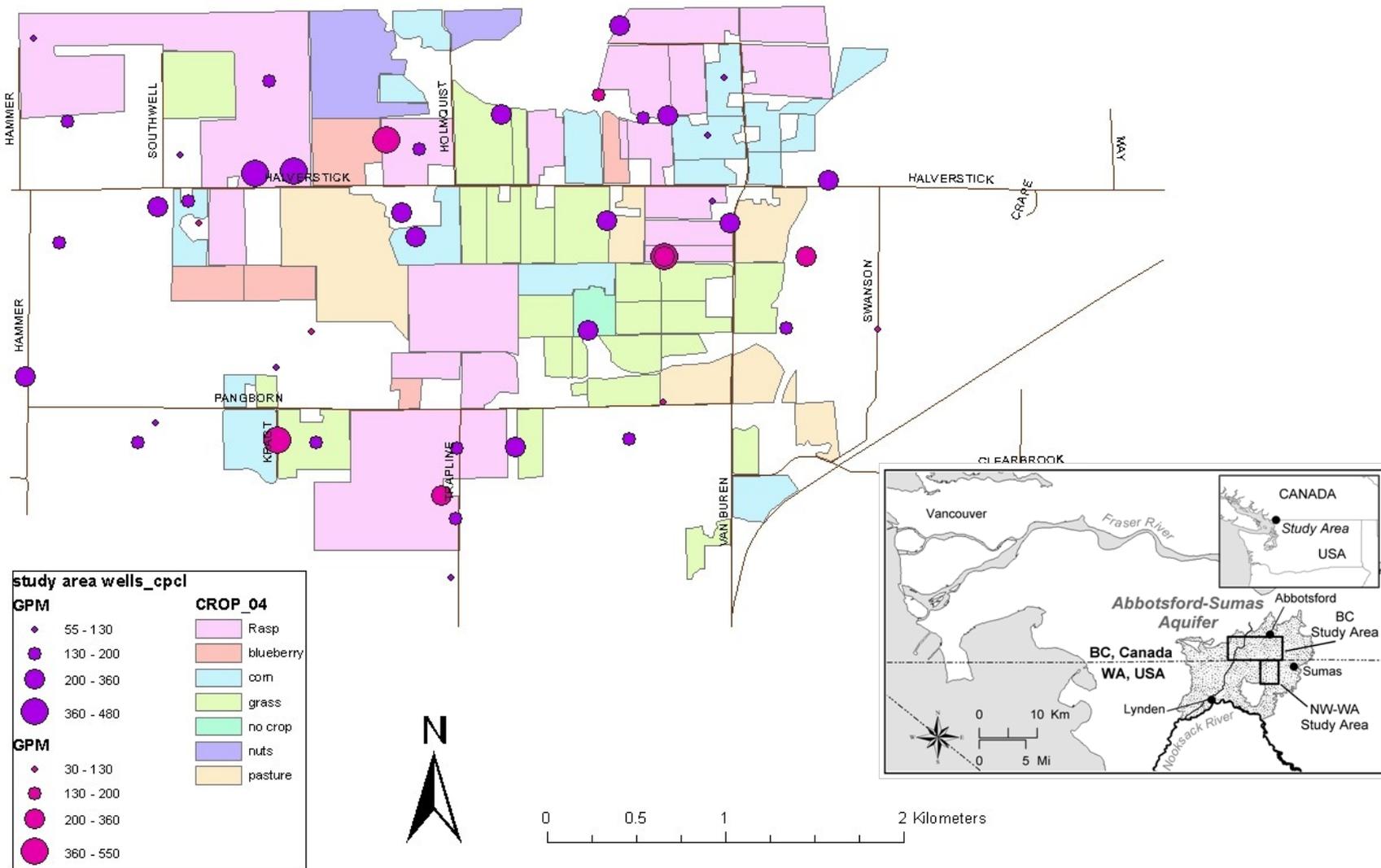


# Groundwater surface water interactions

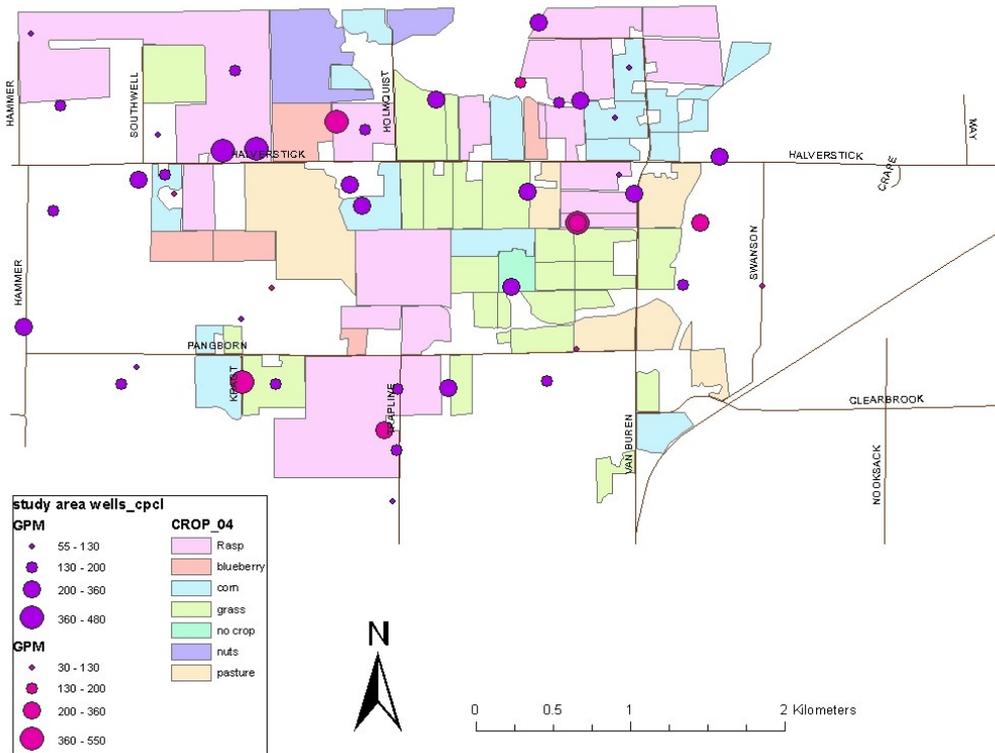
A pumping well can influence streamflow



# Irrigation Wells in the study area



# Sate law highly restricts the water rights of farmers



However, State law allows [Exempt wells](#) (low use wells). So, farmers could sell there land for subdivisions and many homes can drill a well without needing a water right. This is a problem.

